

DELINEATING THE DEVONIAN-MISSISSIPPIAN BOUNDARY BASED ON PALYNOLOGY AT ZUDAÑEZ IN BOLIVIA

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INTRODUCTION

The aim of this contribution is to present the first palynological assemblage recovered from deposits cropping out in the Zudañez area, as part of a major project of Devonian-Permian studies in Bolivia (Fig. 1). The lower stratigraphic deposits in this region characterized by a syncline – anticline complex, correspond to the Devonian Iquiri Formation, represented by heterolithic deposits with micaceous remains and bioturbation. Over an angular

unconformity, an incognitus age-succession starts with alternated gray fangolites and diamicites embedding deformed sand bodies and probable Devonian olistoliths, followed by whitish sandstone beds with pebbly clasts and current structures and red diamicite. Over a parautochthonous contact, another section exhibits whitish sandstone and gray silt/mud and coal levels with plant fossils and interbedded green, gray and red shale/siltstone beds and whitish and red sandstones with current structures, rip-ups, bioturbation and tilloids. The palynological

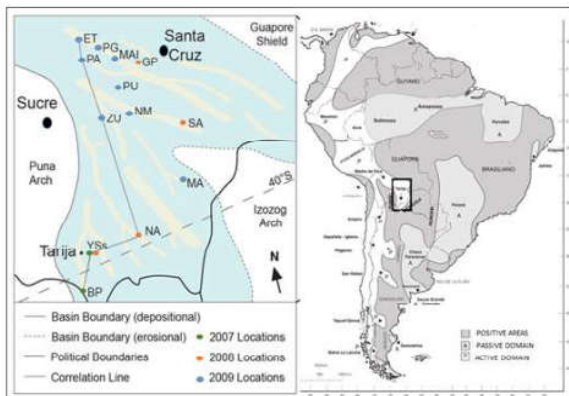


Figure 1 Carboniferous localities surveyed between 2007-2009 (left) in the Tarija-Chaco basin in Bolivia: ET-El Tunal, PG-Pampas Grandes, MAI-Mairana, GP-Ginger Paradise, PA-Pasorapa, PU-Punto Machareti, ZU-Zudañez, NM-Nuevo Mundo, SA-San Antonio de Parapetí, MA-Machareti, NA-Narvaez, VS-Yesera, BP-Balapuca (from Anderson, 2011). Map of South America basins (right, modified from Azucy and di Pasquo, 2000).

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analysis of this succession allowed us to date and establish the stratigraphic units and to propose a correlation to palynofloral zones from South America and elsewhere.

MATERIALS AND METHODS

In the Zudañez area, two samples were obtained from the Iquiri Formation on the way to a place where we collected three more samples of the former and three more from the overlying incognitus deposits. Plant fossils were also recovered from interbedded coal levels. Two more samples were taken from a similar lateral section of the incognitus unit and its uppermost sample was the only one barren. Diamicites were sampled in another location close to the others (Figs. 2-3). A standard palynological method was performed, samples were crushed and treated with hydrochloric acid and after neutralization, attacked with hydrofluoric acid and finally, sieved with 10 µm and 25 µm meshes and slides mounted with glycerine jelly. Palynomorphs were analysed using a light microscope Nikon E200 and illustrated (Plate 1) with a video camera Amuscope 14 Mp. Samples were processed at the Laboratory of Palynostratigraphy and Palaeobotany, catalogued with specific acronym (CICYTP-P) and numbers of the collection housed at CICYTP-CONICET-ER-UADER.

GEOLOGICAL CONTEXT

In Bolivia (Fig. 1), two main deponents are the Madre de Dios Basin in the North extending to Peru and the Tarija-Chaco Basin (Subandean range) in the South with its extension (surface and subsurface) to northern Argentina and western Paraguay. Different stratigraphic names are given to Mississippian units in those deponents (Fig. 2): Cumaná, Kasa and Siripaca formations and Itacua Formation respectively (di Pasquo et al., 2017 and references therein). The Cumaná diamicite beds contain faceted and striated clasts whereas in the Kasa and Siripaca formations prevailed the conglomerate, sandstone, silt and shale lithologies. The Itacua Formation is difficult to differentiate from similar diamicitic units such as Cumaná (?Late Devonian-Tournaisian) and Tarija (early Pennsylvanian) formations. An early Viséan age for the Itacua Formation at Balapuca is confirmed based on a palynological study (di Pasquo, 2007), whereas a latest Devonian age given to another diamicitic deposit at Lajas (close to Santa Cruz, Fig. 1) was put on doubt by Streef et al. (2012). They argued about the chance of having reworked Devonian palynomorphs in agreement with findings documented by di Pasquo and Azucy (1997) and di Pasquo (2003, 2007), among other records. The local unconformable and erosional character of the base of the Cumaná and Itacua formations resulted from erosional and depositional processes related to glacio-marine/

lacustrine environments (Suárez Soruco, 2000; di Pasquo et al., 2017). The change in thicknesses of Mississippian units along with their scarce record are related to the paleogeography of the basin that show a greater structural control. Two phases of uplift are defined during the Chañic Orogeny of Late Devonian to Early Mississippian and Late Mississippian limes in Bolivia and northern Argentina (Tankard et al., 1995; Azucy & di Pasquo, 2000; Suárez Soruco, 2000; Starck & del Papa, 2006). Effects of glacial/deglacial processes are recognized through much of the Mississippian and Early Pennsylvanian succession of the Tarija basin (Starck & del Papa, 2006). Therefore, thicker and widespread diamicite deposits in this region record the local advance and retreat of glaciers into the basin mostly during the Tournaisian and early Viséan and Bashkirian to Kasimovian dated with palynology (e.g. Isaacson et al., 2008; di Pasquo et al., 2017).

PALYNOLOGICAL RESULTS

The five samples from the Iquiri Formation yielded diverse, abundant and well-preserved spores and scarce phytoplankton many bearing pyrite in their exines, and abundant phytoclasts mostly cuticles and tracheids. These assemblages are characterized mainly by the spores *Dibolisporites farraginis*, *Dibolisporites turkulatus*, *Samanisporites triangulatus*, *Grandispora pseudoreticulata*, *Leiotriletes balapucensis*, *Apiculatisporites adavallensis*, *Maranites brasiliensis*, *Hemiruptia tegaulti*, *Quadratisporites* spp., chitinozoans (Fig. 3), being most of them characteristic of the Givetian and Frasnian. Three assemblages were recognized from the incognitus section, in ascending order (Figs. 2-3): Assemblage 1, from the diamicitic section, is composed of indigenous spores (*Anapiculatisporites semicuspidatus*, *Archaeozonitrietes intrastriatus*, *Craspispora invicta*, *C. scrupulosa*, *Cristatisporites indignabundus*, *C. colliculus*, *Cyclogranisporites firmus*, *Dibolisporites microspicatus*, *D. distilicis*, *Grandispora coronata*, *Foveosporites appositus*, *Grandispora maculosa*, *Granulatisporites triconvexus*, *Knovisporites ruhlani*, *Punctatisporites lucidulus*, *Reticulatisporites waloweeiki*, *Vallatisporites cilicaris*, *Velamisporites australiensis*, *Verrucosporites morulatus*, *Waltzisporea lanzonii*), and reworked Devonian spore and phytoplankton species (*Retisporea leptophyta*, *Umbellaspheeridium saharicum*, *Dateriocradus* sp., *Maranites*). Assemblage 2 is characterized by some of the underlying species and the appearance of *Anapiculatisporites concinnus*, *Apiculatisporites caperatus*, *Indotriletes doliandii*, *Reticulatisporites magnificus*, *Tricidatisporites phippisae*, *Waltzisporea polita*, and tetrads of *Verrucosporites* and *Punctatisporites*. Assemblage 3 is less diverse bearing mostly species of the

underlying assemblages, the first record of *Vallatisporites agadensis*, and most abundant spore species of *Punctatisporites* and *Calamospora* and, in a lesser amount, *Verrucosporites* and *Cristatisporites*. Devonian reworked species are scarce in the latter two assemblages. Plate 1

CONCLUSIONS: AGE AND CORRELATION

Most of the mentioned species in Assemblage 1 (Fig. 3) are chronostratigraphically significant for the Tournaisian and Viséan. *Archaeozonitrietes intrastriatus* and *Exallospora coronata* are known in the Viséan of Australia, but not previously recorded in South America (Playford & Melo, 2012; Playford, 2015 and references therein). In Assemblage 2, the appearance of *Anapiculatisporites concinnus*, *Apiculatisporites caperatus*, *Indotriletes doliandii*, *Reticulatisporites magnificus*, *Tricidatisporites phippisae*, *Waltzisporea polita*, confirm a late Viséan age based on the correlation to the Mag Zone Melo & Loboziak and correlative palynofloras of Gondwana, and elsewhere (Fig. 4).

Therefore, the palynological analysis of this succession allowed us to corroborate the presence of late Givetian-early Frasnian deposits of the Iquiri Formation, and to attribute the diamicites to the lower member of the Itacua Formation akin to the late Tournaisian-early Viséan (A1), whereas its upper member to the late Viséan to early Serpukhovian bearing A2 and A3 and plant material. The reworking of late Devonian species into the Itacua Formation confirms active tectonic

processes of Late Devonian and Mississippian as well as the influence of glacial events recognized especially in other South America basins during the Carboniferous (see di Pasquo et al., 2017).

Plate 1. Selected indigenous (1-9) and reworked (10-12) species in Itacua Formation (CICYTP-P) acronym not included, only number of slide, England Finder, size). In 1: *Cristatisporites colliculus* 45-7 Y22(65 µm). In 2: *Anapiculatisporites semicuspidatus* 45-7 W50-148 µm). In 3: *Grandispora maculosa* 45-6 V41(44 µm). In 4: *Reticulatisporites waloweeiki* 71-7 J46(52 µm). In 5: *Tricidatisporites phippisae* 47-6 X48-2(50 µm). In 6: *Exallospora coronata* 42-1 E41(52 µm). In 7: *Archaeozonitrietes intrastriatus* 41-3 Z41-2(55 µm). In 8: *Verrucosporites quasigobbeii* 47-2 Q38(75 µm). In 9: *Reticulatisporites magnificus* 47-3 Y58(100 µm). In 10: *Retisporea leptophyta* 45-5 T41-4(75 µm). In 11: *Umbellaspheeridium saharicum* 41-4 Q27-(110 µm). In 12: *Dateriocradus* sp. 42-3 V46-2(40 µm).

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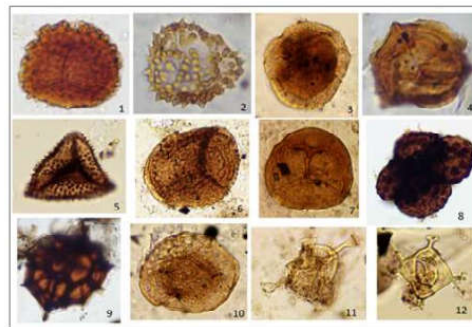


Figure 2 Chronostratigraphic scheme for Bolivia and northern Argentina across the Peru-Bolivia Master Basin.

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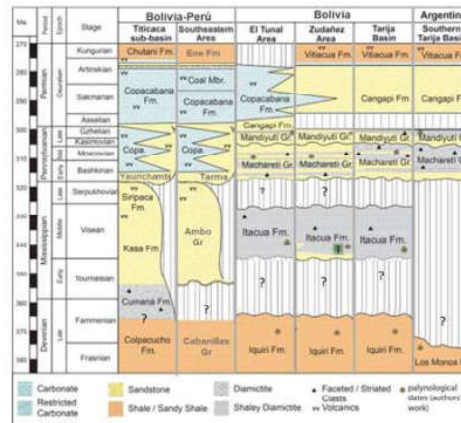


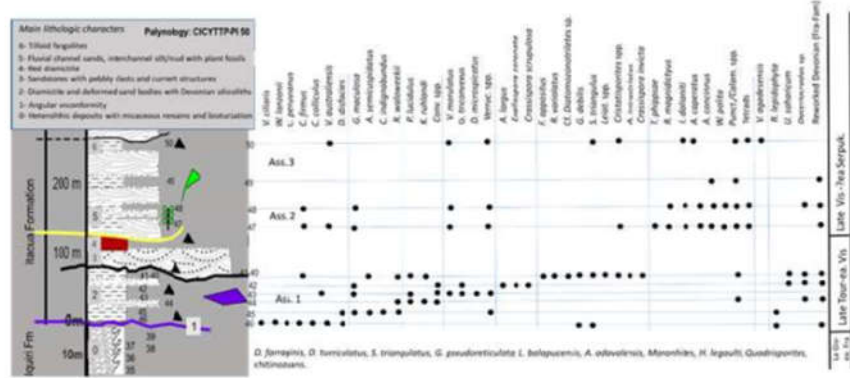
Figure 3 Stratigraphic distribution of selected species at Zudañez (left).

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GEOCHRONOLOGY		REGION	W Europe	N America	E Europe	Australia	Brazil	Argentina	Peru	Bolivia									
PERIOD	EPOCH	STAGE	A	B	C	D	E	F	G	H	I	Zudafñez							
CARBONIFEROUS	MISSISSIPPIAN	SERP.	Amsbergian	TK	P. elegans	No data		R. magnidictyus-V. quasigobbetii (MQ Zone)	Mag Zone	R. magnidictyus Verucosporites sp. 65	No data	No data	No data						
			Pendleian	nitidus-comosus	SM									G. maculosa		R. magnidictyus (Mag Zone)	No data	Assemb. 3	
			Brigantian	verfusus-tracta	AT														No data
		VISEAN	Asbian	nigra-marginatus	stephanophorus	A. largus			?	"Itacua polynofora"	Assemb. 2 (Lycospora)	No data	No data						
			Holartian	TC	D. variabilis D. intermedius K. iteratus C. appendices L. pusillo-M. culta C. multiplocablis									VG exiguus	Cordylotspores-Verucosporites	?	Assemb. 1 (R. waloweekei)	Assemb. 1 (R. waloweekei)	
			Arundian	TS															pusillo-columbaris
		Chadian	pusilla	claviger-macra	decorus-claviger	uncatus	G. spiculiferus	S. preflorus-C. decorus	No data	?	?	No data	No data						
		TOURNAIAN	Ivorian	preflorus-clavata	preflorus-vallatus	P. monotuberculatus								G. spiculiferus	S. preflorus-C. decorus	No data	?	?	No data
			Hastarian	balleatus-polyplycha	?	G. upenab A. septilata	R. arcuatus-W. lanzonii	No data	?	?	No data	No data							
				hibemicus-distinctus	rotatus-explanatus	G. upenab A. septilata							No data						
		verucosus-inchatus	rotatus-explanatus	T. malevkenis															

Figure 4
Biostratigraphic correlation of the assemblages recorded in Zudafñez (see references A-H in di Pasquo and Iannuzzi, 2014). Abbreviations: (W Europe) TS - K. traidatus-K. stephanophorus, TC - P. tessellatus-S. campylosetera, TK - S. thianguis-R. knoxi; (E Europe) VG - M. varonimarginata-V. genitum; (N America) SM - G. spinosa-L. magnificus, AT - S. acadensis-K. traidatus (right).



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