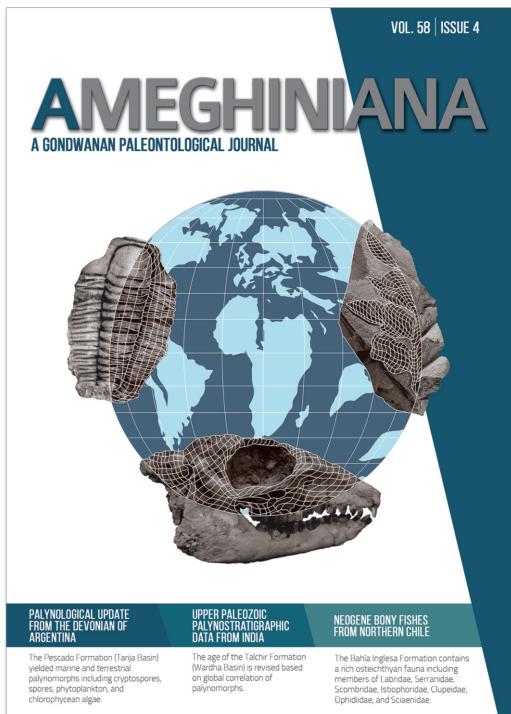




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UPPER PENNSYLVANIAN AGE OF THE LOWER TALCHIR FORMATION IN THE WARDHA BASIN, CENTRAL INDIA, BASED ON GUIDE PALYNOMORPHS PRESENT IN RADIOMETRICALLY-DATED PALYNOZONATIONS IN SOUTH AMERICA, AFRICA, AND AUSTRALIA

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PALYNOLOGICAL UPDATE FROM THE DEVONIAN OF ARGENTINA

The Pescado Formation (Tarija Basin) yielded marine and terrestrial palynomorphs including cryptospores, spores, phytoplankton, and chlorophycean algae.

UPPER PALEOZOIC PALYSTROSTRATIGRAPHIC DATA FROM INDIA

The age of the Talchir Formation (Wardha Basin) is revised based on global correlation of palynomorphs.

NEOGENE BONY FISHES FROM NORTHERN CHILE

The Bahía Inglesa Formation contains a rich osteichthyan fauna including members of Labridae, Serranidae, Scombridae, Istiophoridae, Clupeidae, Ophidiidae, and Sciaenidae.

UPPER PENNSYLVANIAN AGE OF THE LOWER TALCHIR FORMATION IN THE WARDHA BASIN, CENTRAL INDIA, BASED ON GUIDE PALYNOMORPHS PRESENT IN RADIOMETRICALLY-DATED PALYNOZONATIONS IN SOUTH AMERICA, AFRICA, AND AUSTRALIA

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Abstract. The age of the Talchir Formation corresponding to the lower part of the Gondwana (Permo-Carboniferous) Sequence in India is revised in the light of palynostratigraphic data associated with radiometric dating generated in Gondwana. New data was generated from seven samples of the Talchir Formation obtained from Well 131, located in the Penganga area of Wardha Valley Coalfield (Maharashtra State, central India). Two assemblages were demarcated based on the stratigraphic distribution of 63 species of spores (23 taxa), pollen grains (35 taxa), and phytoplankton. Palynoassociation I recognized in the two basal samples yielded few *Punctatisporites* spores and monosaccate pollen grains. Palynoassociation II is demarcated in the overlying three samples characterized by spores and pollen grains that are more diversified and abundant than in the other two samples. Several taxa (*Concavissimisporites grumulus*, *Con verrucosporites confluens*, *Cyclogranisporites gondwanensis*, *Verrucosporites andersonii*, *Latusipollenites quadriscattus*, *Marsupipollenites striatus*, *Pakhapites fusus*, *Striatoabieites multistriatus*, *Vittatina subsaccata*, *Vittatina vittifera*) are shared with palynozones radiometrically constrained to the latest Pennsylvanian–early Cisuralian in Argentina, Brazil, Africa, and Australia. Therefore, we propose a Kasimovian to Ghezelian–Asselian age for the palynoassociations I and II of the Talchir Formation and a correlation with the *Potonieisporites neglectus* and *Plicatipollenites gondwanensis* Zones of Tiwari & Tripathi, respectively. This approach significantly improves the stratigraphic correlations of the Indian palynozones in Gondwana.

Key words. Carboniferous–Permian. Talchir Formation. India. Palynostratigraphy. Radiometric data. Global correlation.

Resumen. PENNSYLVANIANO SUPERIOR, NUEVA EDAD DE LA FORMACIÓN TALCHIR EN LA CUENCA WARDHA, INDIA CENTRAL, BASADA EN PALINOMORFOS GUÍA PRESENTES EN PALYNOZONACIONES RADIOMÉTRICAMENTE DATADAS EN AMÉRICA DEL SUR, ÁFRICA Y AUSTRALIA. La edad de Formación Talchir correspondiente a la parte inferior de la Secuencia Gondwana (Permo-Carbonífero) en India es revisada a la luz de datos palinoestratigráficos asociados a dataciones radiométricas generados en Gondwana. Nuevos datos de siete muestras de la Formación Talchir del Pozo 131 ubicado en el área de Penganga en Wardha Valley Coalfield (Estado de Maharashtra, India central) son generados en este estudio. Dos asociaciones se diferencian con base en la distribución estratigráfica de 63 especies de esporas (23 taxones), granos de polen (35 taxones) y fitoplancton. La Palinoasociación I se reconoció en las dos muestras basales y brindó escasas esporas de *Punctatisporites* y granos de polen monosaccado. La Palinoasociación II se caracteriza en sus tres primeras muestras por esporas y granos de polen más diversificados y abundantes que en las otras dos muestras. En ella aparecen varios taxones (*Concavissimisporites grumulus*, *Con verrucosporites confluens*, *Cyclogranisporites gondwanensis*, *Verrucosporites andersonii*, *Latusipollenites quadriscattus*, *Marsupipollenites striatus*, *Pakhapites fusus*, *Striatoabieites multistriatus*, *Vittatina subsaccata*, *Vittatina vittifera*) documentados en palinozonaciones del Pennsylvaniano–Cisuraliano de Argentina, Brasil, África y Australia. Por ello, proponemos una edad Kasimoviana a Ghezeliano–Asseliana para las palinoasociaciones I y II de la Formación Talchir, y una correlación con las Zonas *Potonieisporites neglectus* y *Plicatipollenites gondwanensis* de Tiwari & Tripathi respectivamente. Este enfoque mejora significativamente las correlaciones estratigráficas de las palinozonas de India en Gondwana.

Palabras clave. Carbonífero–Pérmico. Formación Talchir. India. Palinoestratigrafía. Datos radiométricos. Correlación global.

CHRONOSTRATIGRAPHIC CORRELATIONS of upper Paleozoic successions in Gondwana are predominantly based on palynofloras bearing common palynospecies and especially taking into account those with similar stratigraphic ranges (cosmopolitan) documented across Gondwana. Although, it is proved that first occurrences of many species (**FAD**) are not synchronous reflecting potential paleogeographic patterns of plant migration (e.g., Truswell, 1980; Backhouse, 1991; Lindström, 1995a, 1995b; Playford & Dino, 2000a, 2000b; Stephenson et al., 2003; di Pasquo, 2003, 2009; Souza, 2006; Stephenson, 2008; Vergel, 2008; Modie & Le Hérissé, 2009; di Pasquo et al., 2010, 2015, 2016; Beri et al., 2011; Césari et al., 2011; Barbolini et al., 2016, 2018; Valdez et al., 2020). Moreover, most of the Carboniferous and Permian palynomorphs and plant macrofossils are endemic to Gondwana, preventing the correlation with the International Time Scale (Gradstein et al., 2012; Ogg et al., 2016), which is calibrated using fossils that are not recorded in Gondwanan biotas. For instance, series and stages of the Permian System are named after marine successions standardized with calibrated biotas belonging to the Urals (lower Permian Cisuralian Series, including the Asselian, Sakmarian, Artinskian, and Kungurian stages), southwestern USA (middle Permian Guadalupian Series, including Roadian, Wordian, and Capitanian stages) and southern China (upper Permian Lopingian Series, including the Wuchiapingian and Changhsingian stages). This chronostratigraphic scheme has the approval of the Permian subcommission of the International Commission on Stratigraphy (Jin et al., 1997; Waterhouse, 1997; Archbold, 1999). This schememakes difficult correlations with the vast bulk of Gondwana due to the lack or infrequent records of marine groups such as foraminifers, conodonts, and ammonoids of the tropical marine international stratotypes (Metcalfe et al., 2017; Henderson, 2018). Late Paleozoic sequences of western Australia and northern Amazonas and Parnaíba basins of Brazil bearing the richest marine faunas (Picarelli et al., 1991; Archbold, 1999; Metcalfe & Sone, 2008) offer the best opportunity to correlate with some precision to those international marine stratotypes, but are absent or extremely rare in most sequences of southern Gondwana. Nevertheless, correlations between farther regions in Gondwana are still possible due to the existence of index taxa having confident FADs, enhanced by absolute dates when interbedded igneous rocks are available in the

successions (e.g., Bangert et al., 1999; Santos et al., 2006; Guerra-Sommer et al., 2008; Stephenson, 2009; Gulbranson et al., 2010; Césari et al., 2011; Mori et al., 2012; di Pasquo et al., 2015; Barbolini et al., 2016, 2018; Griffis et al., 2019; Valdez et al., 2020).

In most of the Indian Permian, with a few exceptions (Tiwari & Tripathi, 1992), palynological zones are defined using the dominance datum of taxa (e.g., Aggarwal & Jha, 2013) at generic level, which prevents accurate comparisons and correlations with FAD zonal schemes elsewhere in Gondwana. Additionally, a taxonomic revision of illustrated and listed endemic species of the Permian successions in India (e.g., Tiwari & Tripathi, 1992; Aggarwal & Jha, 2013, and their references) is required as previously highlighted in several publications (Truswell, 1980; Stephenson, 2008; Modie & Le Hérissé, 2009; di Pasquo et al., 2015, 2021; Barbolini et al., 2018). In the case of those taxa that are not illustrated, a revision of samples of their respective works or the development of new studies to identify Pennsylvanian and Permian guide species is mandatory (see Murthy et al., 2020; di Pasquo et al., 2021). Murthy et al. (2020) obtained a palynoassemblage attributed to the *Potonieisporites neglectus* Zone of Tiwari & Tripathi (1992) from the Lower Talchir Formation that appears below the *Plicatipollenites gondwanensis* Assemblage Zone bearing diagnostic Permian spores/pollen grains. An undoubted late Carboniferous age is assigned to this basal Talchir Formation by Murthy et al. (2020), confirming the interpretation of Mukhopadhyay et al. (2010) and few previous studies (Casshyap & Qidwai, 1974; De, 1979; Raja Rao, 1982, 1983, 1987). The early Permian age given by other researchers due to the absence of characteristic late Carboniferous fossils or palynofloras likely by erosion or non-deposition of older strata (Tiwari & Tripathi, 1988, 1992; Venkatachala & Tiwari, 1988; Tiwari, 1996; Vijaya, 1996; Jha et al., 2018) is discarded. Therefore, despite the current lack of absolute dates from the late Paleozoic Indian basins, more accurate palynotaxonomic studies are mandatory to attempt more accurate comparisons and correlations with other palynofloras across Gondwana. We present the results of the palynologic study of core samples from the Talchir Formation at Penganga in Wardha Basin, Maharashtra State, Central India (Fig. 1). The age of the assemblages is based on the recognition of diagnostic species and correlations established with palynozones radiometrically

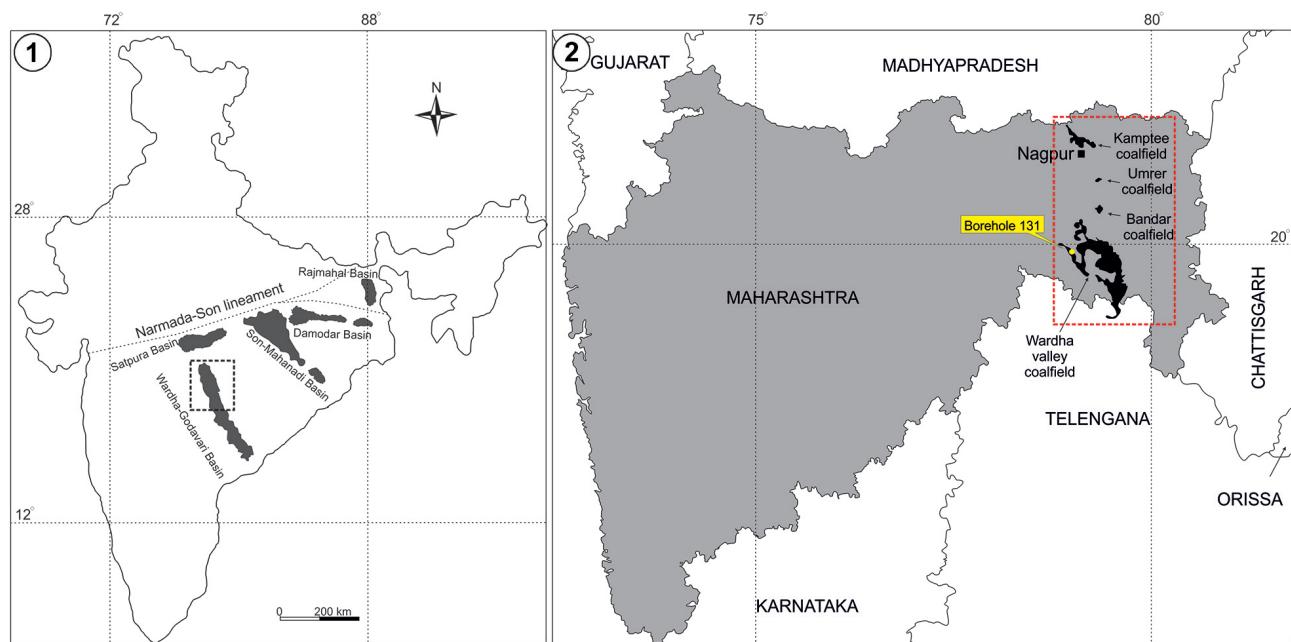


Figure 1. 1, Map showing the Gondwana basins in India (after Geological Survey of India). 2, Map of Maharashtra State showing the coalfields in the Wardha Basin and the location of Borehole 131 of the present study (after Kavali *et al.*, 2016).

constrained elsewhere in Gondwana. A better age constraint is ongoing for this unit in other basins of India based on the present work.

PREVIOUS WORKS

A review of the voluminous literature from different Gondwana basins of India reveals that palynological studies of the Talchir Formation are relatively scarce compared to the plethora of data generated from the overlying coal-bearing lower Permian Barakar and upper Permian Raniganj formations (and its equivalents Kamthi and Bijori) due to the extensive coal mining activity (Mukhopadhyay *et al.*, 2010). Another possible factor is the scarcity of drill cores of the Talchir sediments available for scientific studies because the exuberant costs prevent the continuity of the drilling below the last coal-bearing sediments just over that unit. However, few palynological works were possible in cases where Talchir sediments were unexpectedly intersected in faulted horizons in due course of drilling. Few studies were also carried out on outcrops.

A literature survey of the palynological works from Talchir sediments of the Wardha Basin is relatively less and requires further studies. Many palynologic studies were carried out in the lower Permian Barakar and upper Permian

Kamthi formations (Agashe & Chitnis, 1970, 1972a, 1972b; Anand-Prakash & Khare, 1974; Bharadwaj & Anand-Prakash, 1974; Agashe *et al.*, 1981; Sarate, 1985; Bhattacharrya, 1997; Kumar & Jha, 2000; Jha *et al.*, 2007; Sabina *et al.*, 2007; Shivanna *et al.*, 2008, 2011, 2014; Kalkar *et al.*, 2010; Kavali *et al.*, 2010, 2016; Jha *et al.*, 2011; Murthy & Sarate, 2016; Sarate *et al.*, 2016, 2017; Murthy *et al.*, 2017, 2019). Other studies were on megaspores (Agashe, 1979; Tewari *et al.*, 2004; Murthy *et al.*, 2017) and macrofloras (Bunbury, 1861; Feistmantel, 1881; Chitnis & Vagyan, 1979; Varadpande, 1977a, 1977b; Chandra & Prasad, 1981; Raja Rao, 1982; Tewari & Rajanikanth, 2001; Singh *et al.*, 2005; Tewari, 2007, 2008). Fewer studies concern seeds (Sundaram & Nandi, 1984; Tewari, 2007; Tewari *et al.*, 2012) and gymnospermous woods (Agashe & Prasad, 1989; Agashe & Kumar, 1996, 2001; Agashe, 2001; Agarwal *et al.*, 2007).

Hitherto there is only one published palynological study from the Talchir deposits of the Wardha Basin by Lele (1979). He discovered for the first time a definite basal Talchir microflora from the Penganga river section, close to the glacial pavement near Irai, on the east bank of the Penganga river about 1.6 km to the south of its confluence with the Wardha river. There, the Talchir Formation lies unconformably on the limestones of the Precambrian Penganga Formation, the

surface of which exposes the classical Penganga Pavement with grooves and striations studied by Smith (1963), who established it as unequivocal evidence of the late Paleozoic Gondwana glaciation (LPIA). The palynoassemblage recovered from the basal Talchir Formation contains 70% acritarchs and 30% miospores dominated by radial monosaccate pollen grains, notably of the genera *Potonieisporites*, *Cannanoropolis*, and *Plicatipollenites*.

Our work is the second palynofloral report from the Talchir Formation of Wardha Basin, and a taxonomic updated list of the species provided by Lele (1979) is included, along with a discussion on its age.

GEOLOGY OF THE AREA

The Gondwana deposits in India are exposed and spread over an enormous extent of 50,000 km² and collectively represent about 6,000 m of strata (Mukhopadhyay *et al.*, 2010). The Gondwana basins of Peninsular India show certain specific characteristics in terms of their distribution pattern and mode of evolution. These occur as conspicuous rectilinear belts of basins of failed rift origin (Acharyya, 2000). These basins were embryonic in dimension during the initial stage of Gondwana deposition (*i.e.*, during the Talchir Period), where sedimentation commenced with deposition of glacial and glaciogene fluvio-lacustrine sediments. Evidently, the growth of these deposits continued through a regional tensional regime, in which the basins were characterized by extensional faulting parallel to the rift trend creating half-grabens that formed different sub-basins within the master basin belt (Acharyya, 2000; Tewari & Mahejima, 2010). The evolution and pattern of these Permo-Cretaceous sedimentary basins along particular linear belts are genetically related to the ancient seismo-tectono-thermal activity belts, which particularly became active during the late Paleozoic. The development of basins fits with the structural geometry of the basement modeled by distensional events that resulted in rectilinear distribution patterns (Chakraborty, 1999). The Indian shield constitutes a multi-cratonic mosaic of several Precambrian crustal blocks assembled together along prominent tectonic joints and mobile belts through the incidence of recurrent tectono-thermal events from Archaean to Late Proterozoic (Chakraborty, 1999). Four major belts along which these basins are distributed, viz., (a) Satpura-Son-Damodar belt

(E-W), (b) Godavari-Wardha belt (NNW–SSE), (c) Mahanadi belt (NNW–SSE), and (d) Rajmahal-Birbhum belt (N–S) (Fig. 1.1).

The Wardha Basin holds a premier position in India for having a considerable share of reserve of thermal grades non-coking coal for catering to the demand of coal in the western part of the country. The late Paleozoic Gondwana outcrops in the Wardha Basin are concentrated in the north-eastern parts of the Maharashtra State, bordering the states of Madhya Pradesh and Telangana (Fig. 1.2). There are several coalfields identified in the Wardha Basin and are concentrated in the districts of Nagpur (Umred Colliery, Sillewara Colliery, Kamptee Colliery, Bina, and Ghatochan), Chandrapur (Wardha Valley, Ghugus, Ballarpur, Hindustan Lalpeth Colliery, Rajura, Tilwasa, and Sasi Colliery) and Wardha (Majri, Warora, Bandar) (Agashe & Chitnis, 1970; Raja Rao, 1987).

The Wardha Valley Coalfield is located in the southeastern part of Maharashtra State in the districts of Yavatmal and Chandrapur. An area of 4,000 km² is estimated, between 19° 30' and 20° 27' N, and 78° 50' and 79° 49' E (Raja Rao, 1987). The NNW–SSE axis corresponds to the strike of the coal-bearing successions and is around 100 km long with a maximum width of about 80 km. The extension of this coalfield in north-northwest direction beneath the Deccan Trap is not discarded. A south-south eastward continuation of the Wardha Valley Coalfield in Godavari Valley Coalfield of the Telangana State (previously included in Andhra Pradesh State) is considered. The area comprises Precambrian and late Paleozoic Gondwana successions and Deccan Trap rocks covered with black soil and alluvium. The lithostratigraphic sequence in the Wardha Valley Coalfield is given in Figure 2. The general slope of the area is towards the south and drained mainly by the Wardha, Penganga, and Erai rivers. North-eastern area is drained by the Erai river and its tributaries, whereas the southern part of the area is drained by Penganga flowing along the south boundary of the coalfield.

The regional structure of the Wardha Valley Coalfield is a broad anticline plunging towards NNW. Both, the western and eastern limbs of this anticline have been proved to be coal-bearing. The western limb dips westerly, whereas the eastern limb dips easterly. The Gondwana sedimentation in Wardha Valley Coalfield has taken place in NW–SE trending

Age	Formation	Lithology
upper Permian to lower Triassic	Kamthi	Yellow to brown fine to coarse grained sandstones, shale and variegated clays
UNCONFORMITY		
middle Permian	Motur	Medium to fine to grained variegated sandstones, variegated clays and shale
lower Permian	Barakar	Grey to white fine to coarse grained sandstones, thin clay bands, shale, intercalation of shale & sandstone, sandy shale, shaly sandstone, carbonaceous shale, shaly coal and coal.
upper Carboniferous to lower Permian	Talchir	Green shale, grey shale, siltstones and sandstones
UNCONFORMITY		
Precambrian	Vindhyan Formation	Limestone

Figure 2. Lithostratigraphic succession in the Wardha Valley Coalfield (unpublished report of Central Mine Planning and Design Institute Limited).

rift basins separated by Vindhyan, Deccan Trap/Lameta and Kamthi formations unconformably overlie the Motur and coal-bearing Barakar formations, preventing the coal seams outcrop on the surface. NW–SE trending normal faults are major structural features of Wardha Valley Coalfield. These strike faults have caused the repetition of strata in many parts of the coalfield. As a result, the coal seams have occurred at shallow depths in many areas of the Wardha Valley Coalfield, opening additional possibilities of opencast mining in this coalfield.

The coalfield has elliptically aligned coal prospects within the Barakar Formation around the core of Talchir Formation, which occupies its central part. The Gondwana deposits boundary toward the east is delineated by the exposures of Archaean rocks. The intervening area between the faults outlines of coal-bearing deposits towards the northeast, southwest, and southeast is occupied by the Vindhyan Formation, whereas in its central part is the Talchir Formation. The Deccan Trap covers the Gondwana sediments towards the west and north. Strata of the Lameta Formation are available in patches towards the north. The coal-bearing Barakar Formation is exposed only in the western part of the coalfield in isolated patches covered by the overlying Kamthi Formation.

MATERIALS AND METHODS

Nine core samples were collected from Borehole 131 drilled in Penganga area of Wardha Valley Coalfield in Wardha Basin, Central India. The borehole lies between 19° 49' 4.43" N and 79° 8' 3.59" E (Fig. 1.2). The depth of the borehole is 66 m, and the Talchir Formation comprises grey greenish shales and siltstones that lie over the limestones of the Precambrian Vindhyan basement and overlain by the sandstones of the Barakar Formation (Fig. 3). Samples were processed applying Hydrochloric acid (HCl) 10% and Hydrofluoric acid (HF) 40%, washed with distilled water, sieved through a 10 µm mesh, and slides prepared with +10 residues and Canada balsam. The residues were divided into two sets, and one was oxidized with Nitric acid (HNO₃) 40% for two hours and Potassium hydroxide (KOH) 10% applied for two minutes (see di Pasquo *et al.*, 2021). Quali-quantitative (200 palynomorphs per sample) examination was performed using a Leica DM 2500 light microscope, and pictures of specimens were captured with a Leica DFC290 camera using LAS V4.6 software. Materials were processed at the laboratory of the Birbal Sahni Institute of Palaeosciences (**BSIP**) and housed under the acronyms BSIP 16598–16630. Figured specimens are indicated by the respective acronym followed by the number of the slide and the England Finder coordinate.

RESULTS

Characteristics of the palynoassemblages

The basal grey shale samples, S.1 and S.2, lying over the limestones of the Precambrian Vindhyan basement at a depth of 66 and 64 m, contain poorer assemblages and are the least diversified, represented by monosaccate grains. The overlying greenish shale samples, S.3–S.5, between the depth-interval of 62–57 m, contain the most diversified and best-preserved palynomorphs. The samples S.6–S.7 from the siltstones of the Talchir Formation were poor in their yield. The sandstone samples, S.8 and S.9, of the overlying Barakar Formation were unproductive. The relative frequency and vertical distribution of 63 species composed of spores (23 taxa), pollen grains (35 taxa), and phytoplankton support the definition of two assemblages (Tab. 1), and significant forms are illustrated (Figs. 4–7). The analysis of the chrono-biostratigraphic documentation of spores and pollen grains reveals diagnostic species that allow us to establish correlations addressed in discussion (Fig. 8; Tab. 2).

Palynoassemblage I

It is recognized in the basal-most grey shale sample S.1 at 66 m and S.2 at 64 m depth (Fig. 3), characterized by poorly diverse spores of pteridophytes (*Punctatisporites gretensis*) and cordaitean and coniferalean monosaccate pollen species of *Cannanoropollis*, *Plicatipollenites*, and *Potonieisporites* (Tab. 1).

Palynoassemblage II

The overlying greenish shale samples between the depth intervals of 62 m (S.3)–57 m (S.5) yielded abundant and diversified palynomorphs fairly well preserved. The first appearance of several Permian marker taxa such as the trilete spore *Concavissimisporites grumulus*, *Converrucosporites confluens*, *Verrucosporites andersonii* (Pteridophyta), and gymnospermous pollen grains *Latusipollenites quadrisaccatus*, *Pakhapites fusus*, *Striatoabietites multistriatus*, *Vittatina subsaccata*, and *Vittatina vittifera* is documented. Frequent or dominant species of *Plicatipollenites*, *Cannanoropollis*, and *Potonieisporites*, as well as other groups of pollen grains widely distributed in Gondwana, mainly with Pennsylvanian–

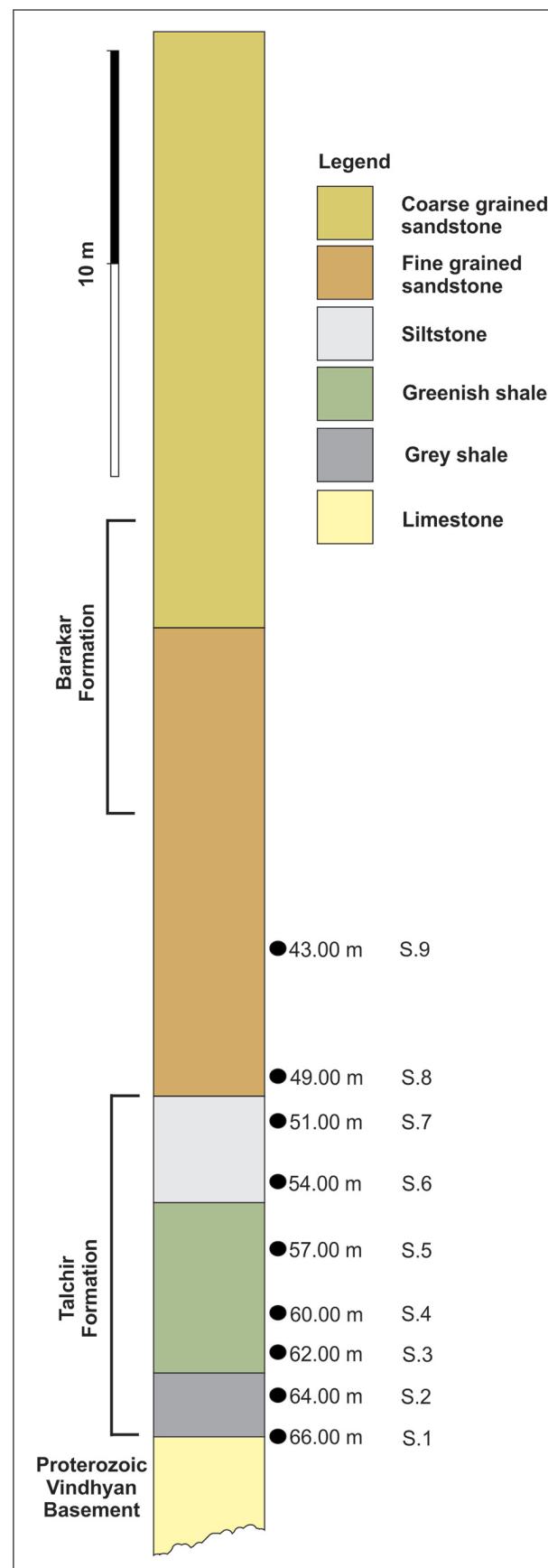


Figure 3. Lithology of Borehole 131 showing the position of samples studied.

TABLE 1. Distribution of taxa and demarcation of Palynoassemblage I and Palynoassemblage II of Talchir Formation in Borehole 131, Wharda Basin.

Taxa			Assemblage							
Spores	/	Samples	I			II			*Lele (1979)	
			1	2	3	4	5	6		
<i>Brevitriletes cornutus</i> (Balme & Hennelly, 1956) Backhouse, 1991					P	P	P			
<i>Calamospora hartungiana</i> Schopf in Schopf, Wilson & Bentall, 1944				F	F	A				
<i>Calamospora landiana</i> Balme, 1970				F	F	F		X		
<i>Calamospora liquida</i> Kosanke, 1950				F	F	F		X		
<i>Concavissimisporites grumulus</i> Foster, 1979						P				
<i>Converrucosporites micronodosus</i> (Balme & Hennelly, 1956) Playford & Dino, 2002				F	P	P				
<i>Converrucosporites confluens</i> (Archangelsky & Gamerro, 1979) Playford & Dino, 2002				P						
<i>Convolutispora ordonensis</i> Archangelsky & Gamerro, 1979				P						
<i>Cristatisporites conatus</i> (Lele & Makada, 1972) comb. nov.						P		X		
<i>Cristatisporites pseudozonatus</i> (Lele & Makada, 1972) Jones & Truswell, 1992				P		P				
<i>Cristatisporites</i> sp.						P		X		
<i>Cyclogrananisporites gondwanensis</i> Bharadwaj & Salujha, 1964								X		
<i>Granulatisporites austroamericanus</i> Archangelsky & Gamerro, 1979 (closely similar to <i>Microbaculispora tentula</i> Tiwari, 1965)				P			X			
<i>Grossusporites</i> sp.				P						
<i>Horriditriletes gondwanensis</i> (Tiwari & Moiz, 1971) Foster, 1975				P	P	P				
<i>Horriditriletes ramosus</i> (Balme & Hennelly, 1956) Bharadwaj & Salujha, 1964				F		P				
<i>Horriditriletes uruguaiensis</i> (Marques-Toigo, 1974) Archangelsky & Gamerro, 1979				P						
<i>Leiotriletes directus</i> Balme & Hennelly, 1956				P	P	F				
<i>Leiotriletes virkii</i> Tiwari, 1965				F	F	F				
<i>Lophotriletes rectus</i> Bharadwaj & Salujha, 1964					P					
<i>Punctatisporites gretensis</i> Balme & Hennelly, 1956			X	X	F	F	A	X	sp.	
<i>Verrucosporites verrucosus</i> (Ibrahim, 1932) Ibrahim, 1933				P	P					
<i>Verrucosporites andersonii</i> (Anderson, 1977) Backhouse, 1988				P	P					
Pollen grains	/	Samples	1	2	3	4	5	6	7	*Lele (1979)
<i>Caheniasaccites flavatus</i> (Bose & Kar, 1966) Azcuy & di Pasquo, 2000							P		X	
<i>Cannanaropollis densus</i> (Lele, 1964) Bose & Maheshwari, 1968				X	F	P	P	X		X
<i>Cannanaropollis janakii</i> Potonié & Sah, 1960			X	X	F	F	P	X	X	X
<i>Cannanaropollis mehtaee</i> (Lele, 1964) Bose & Maheshwari, 1968			X	X	F					
<i>Cannanaropollis triangularis</i> (Mehta, 1944) Bose & Maheshwari, 1968			X		P					
<i>Cannanaropollis trigonalis</i> (Bose & Maheshwari, 1966) Bose & Maheshwari, 1968				P		P				
<i>Costatascyclus crenatus</i> (Felix & Burbridge, 1967) Urban, 1971				P						
<i>Cycadopites cymbatus</i> (Balme & Hennelly, 1956) Segroves, 1970			X	X	A	A	F		X	
<i>Divarisaccus lelei</i> Venkatachala & Kar, 1966				P	P					
<i>Illinites talchirensis</i> (Lele & Makada, 1972) Azcuy, di Pasquo & Ampuero, 2002							P			

TABLE 1. Continuation.

Taxa			Assemblage							
Pollen grains	/	Samples	1	2	3	4	5	6	7	*Lele (1979)
<i>Latusipollenites quadrisaccatus</i> Marques-Toigo, 1974					P		P			
<i>Limitisporites rectus</i> Leschik, 1956									X	
<i>Limitisporites hexagonalis</i> Bose & Maheshwari, 1968						P		X		
<i>Lunatisporites noviaulensis</i> Foster, 1979								X		
<i>Lunatisporites varisectus</i> Archangelsky & Gamerro, 1979					P					
<i>Marsupipollenites striatus</i> (Balme & Hennelly, 1956) Foster, 1975					P		P			
<i>Marsupipollenites triradiatus</i> Balme & Hennelly, 1956					F	A	A		X	
<i>Pakahpites fusus</i> (Bose & Kar, 1966) Menéndez, 1971					P					
<i>Plicatipollenites gondwanensis</i> (Balme & Hennelly, 1956) Lele, 1964			X	X	F	F	F	X	X	X
<i>Plicatipollenites malabarensis</i> (Potonié & Sah, 1960) Foster, 1979					A	A	F	X		X
<i>Plicatipollenites trigonalis</i> Lele, 1964			X	X	F	A	F	X	X	X
<i>Potonieisporites barrelis</i> Tiwari, 1965				X			P			X
<i>Potonieisporites congoensis</i> Bose & Maheshwari, 1968			X			P	P			
<i>Potonieisporites densus</i> Maheshwari, 1967				X						X
<i>Potonieisporites lelei</i> Maheshwari, 1967			X	X	F		P	X	X	X
<i>Potonieisporites magnus</i> Lele & Karim, 1971							P			
<i>Potonieisporites neglectus</i> Potonié & Lele, 1961			X	X	P					X
<i>Potonieisporites novicus</i> (Bharadwaj, 1954) Poort & Veld, 1997				X		P	P	X		
<i>Protohaploxylinus latissimus</i> (Luber in Luber & Waltz, 1941) Samoilovich, 1953					P					
<i>Protohaploxylinus limpidus</i> (Balme & Hennelly, 1956) Balme & Playford, 1967					P	F	P			
<i>Protohaploxylinus perfectus</i> (Naumova, 1939) Samoilovich, 1953					P		P			
<i>Scheuringipollenites maximus</i> (Hart, 1960) Tiwari, 1973									X	
<i>Striatoabieites multistriatus</i> Foster, 1979					P		P			
<i>Vittatina subsaccata</i> Samoilovich, 1953							P			
<i>Vittatina vittifera</i> (Luber in Luber & Waltz, 1941) Samoilovich, 1953							P			
Algae	/	Samples	1	2	3	4	5	6	7	*Lele (1979)
<i>Botryococcus braunii</i> Kützing, 1849			X	X	A	A	F		X	X
<i>Brazilea scissa</i> (Balme & Hennelly, 1956) Foster, 1975					P		P			
<i>Cymatiosphaera gondwanensis</i> (Tiwari, 1965) Backhouse, 1991					P		P			
<i>Tetraporina punctata</i> (Tiwari & Navale, 1968) Kar & Bose, 1976					P		P			

The frequency of species expressed as: X= presence (not counted in the samples at the depths of 66 m, 64 m, 54 m, 51 m), P (present)= < 2.5 %, F (frequent)= 2.5–7 %, A (abundant)= > 7 %. * For synonymy of species illustrated by Lele (1979) proposed herein see in the text under taxonomy.

Permian ranges, include *Marsupipollenites triradiatus*, *Marsupipollenites striatus* (Pteridospermophyta), *Cycadopites cymbatus* (Cycadophyta/Ginkgophyta). Among spores, sphenophytes (*Calamospora hartungiana*, *Calamospora landiana*, *Calamospora liquida*), pteridophytes (*Brevitriletes cornutus*, *Convolutispora ordonensis*, *Converrucosporites micronodosus*, *Horriditriletes ramosus*, *Horriditriletes uruguayensis*, *Horriditriletes gondwanensis*, *Punctatisporites gretensis*, *Leiotriletes directus*, *Leiotriletes virkii*) and few lycophytes (*Cristatisporites pseudozonatus*, *Cristatisporites conatus*) are present.

SYSTEMATIC PALEONTOLOGY

Suprasubturma LAMINATITRILETES

Smith & Butterworth, 1967

Subturma ZONOLAMINATITRILETES

Smith & Butterworth, 1967

Infraturma CINGULICAVATI

Smith & Butterworth, 1967

Genus *Cristatisporites* (Potonié & Kremp, 1954)

Butterworth, Jansonius, Smith & Staplin, 1964

Type species. *Cristatisporites indignabundus* Loose in Potonié & Kremp, 1954. Pennsylvanian of Germany.

Cristatisporites conatus

(Lele & Makada, 1972) comb. nov.

Basionym. *Jayantisporites conatus* Lele & Makada, 1972.

Remarks. The general morphology is similar to *Cristatisporites menendezii* (Menéndez & Azcuy, 1972) Playford, 1978 emend. Césari, 1986 and *Cristatisporites stellatus* (Azcuy, 1975) Gutiérrez & Limarino, 2001, although it differs in having a narrower incomplete cingulizone composed of cristate ornamentation. *Cristatisporites pseudozonatus* (Lele & Makada, 1972) Jones & Truswell, 1992 is characterized by a larger and denser cristate ornamentation. Backhouse (1988) considered *Cristatisporites inconstans* Archangelsky & Gamerro, 1979 a junior synonym of the latter taxon, whereas we agree with other authors that maintain both species separated based on minor differences of their ornamentation and completeness of cingulizones (see Jones &

Truswell, 1992; Césari *et al.*, 2019; Backhouse & Mory, 2020).

On the other hand, we propose a re-assignment of the illustrated species from Talchir Formation published by Lele (1979) as follows: *Callumispora uniformis* (Tiwari, 1968) Lele & Chandra, 1977 (Lele, 1979, pl. 1.1) = *Punctatisporites gracilis* Anderson, 1977; *Plicatipollenites indicus* Lele, 1964 (Lele, 1979, pl. 1.2), and *Potonieisporites crassus* Lele & Chandra, 1977 (Lele, 1979, pl. 1.5) = *Plicatipollenites malabarensis* (Potonié & Sah, 1960) Foster, 1975; *Plicatipollenites densus* Lele, 1964 (Lele, 1979, pl. 1.4) = *Plicatipollenites gondwanensis* (Balme & Hennelly, 1956) Lele, 1964; *Potonieisporites crassus* Lele & Chandra, 1977 (Lele, 1979, pl. 1.6) = *Potonieisporites densus* Maheshwari, 1967; *Potonieisporites jayantensis* Lele & Karim, 1972 (Lele, 1979, pl. 1.7) = *Potonieisporites barrelis* Tiwari, 1965 (see Azcuy & di Pasquo, 2000); *Parasaccites densicorpus* Lele, 1964 (Lele, 1979, pl. 1.10–1.11) and *Parasaccites densus* Maheshwari, 1967 (Lele, 1979, pl. 1.12) = *Cannanoropollis densus* (Lele, 1964) Bose & Maheshwari, 1968; *Parasaccites korbaensis* Bharadwaj & Tiwari, 1964 (Lele, 1979, pl. 1.13), *Parasaccites diffusus* Tiwari, 1965 (Lele, 1979, pl. 1.14); and *Parasaccites fimbriatus* Maheshwari, 1969 (Lele, 1979, pl. 1.15) = *Cannanoropollis janakii* Potonié & Sah, 1960 (see Azcuy & di Pasquo, 2000); *Caheniasaccites densus* Lele & Karim, 1972 (Lele, 1979, pl. 1.16) = *Caheniasaccites flavatus* Bose & Kar, 1966 (see Azcuy & di Pasquo, 2000); *Gondwanopollis densus* Lele & Chandra, 1977 (Lele, 1979, pl. 1.17) = *Potonieisporites lelei* Maheshwari, 1967 (see Tab. 1). The specimen cf. *Crescentipollenites* (Lele, 1979, pl. 1.18) would fit in *Lunatisporites* sp. based on the description of the striated cappa.

DISCUSSION

Age of the assemblages

Most of the Gondwana palynostratigraphic works published in India applied the palynozonation scheme proposed by Tiwari and Tripathi (1992). They have demarcated three palynozones in the basal Talchir Formation: *Potonieisporites neglectus*, *Plicatipollenites gondwanensis*, and *Cannanoropollis janakii* (= *Parasaccites korbaensis*) in stratigraphic order constrained to the early Permian based on the general dominance of monosaccates and the associated occurrence of the marine *Eurydesma* fauna. Murthy *et al.* (2020) revised the age of these palynozones based on a critical analysis of the composition of these assemblage zones, their reference

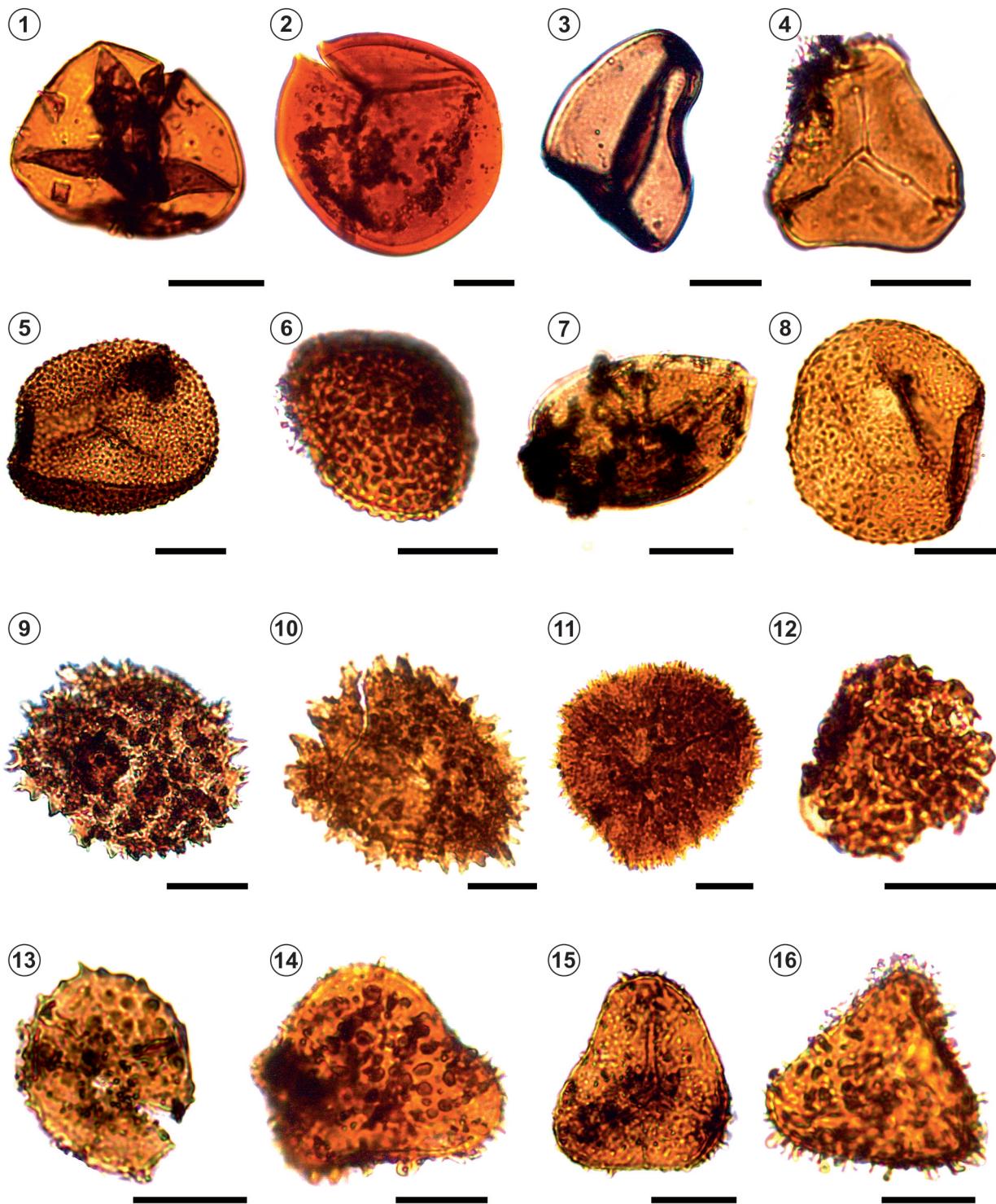


Figure 4. 1, *Calamospora landiana* Balme, 1970, BSIP 16620, EF: M.45; 2, *Punctatisporites gretensis* Balme & Hennelly, 1956, BSIP 16604, EF: J.49; 3, *Leiotriletes directus* Balme & Hennelly, 1956, BSIP 16605, EF: P.52; 4, *Leiotriletes virkii* Tiwari, 1965, BSIP 16612, EF: L.34; 5, *Cyclogranisporites gondwanensis* Bharadwaj & Saluja, 1964, BSIP 16628, EF: K.57; 6, *Verrucosporites verrucosus* (Ibrahim, 1932) Ibrahim, 1933, BSIP 16618, EF: J.43.4; 7, *Grossusporites* sp., BSIP 16620, EF: X.35; 8, *Verrucosporites andersonii* (Anderson, 1977) Backhouse, 1988, BSIP 16598, EF: G.53.3; 9, *Cristatisporites conatus* (Lele & Makada, 1972) comb. nov., BSIP 16626, EF: V.50; 10, *Cristatisporites pseudozonatus* (Lele & Makada, 1972) Jones & Truswell, 1992, BSIP 16615, EF: T.46.4; 11, *Cristatisporites* sp., BSIP 16625, EF: O.44.2; 12, *Convolutispora ordonensis* Archangelsky & Gamerro, 1979, BSIP 16612, EF: V.50; 13, *Brevitriletes cornutus* (Balme & Hennelly, 1956) Backhouse, 1991, BSIP 16618, EF: W.51; 14, *Horriditriletes ramosus* (Balme & Hennelly, 1956) Bharadwaj & Saluja, 1964, BSIP 16604, EF: H.67; 15, *Horriditriletes gondwanensis* (Tiwari & Moiz, 1971) Foster, 1975, BSIP 16624, EF: M.28.1; 16, *Horriditriletes uruguayensis* (Marques-Toigo, 1974) Archangelsky & Gamerro, 1979, BSIP 16618, EF: W.41.3. Scale bars equal 20 μ m.

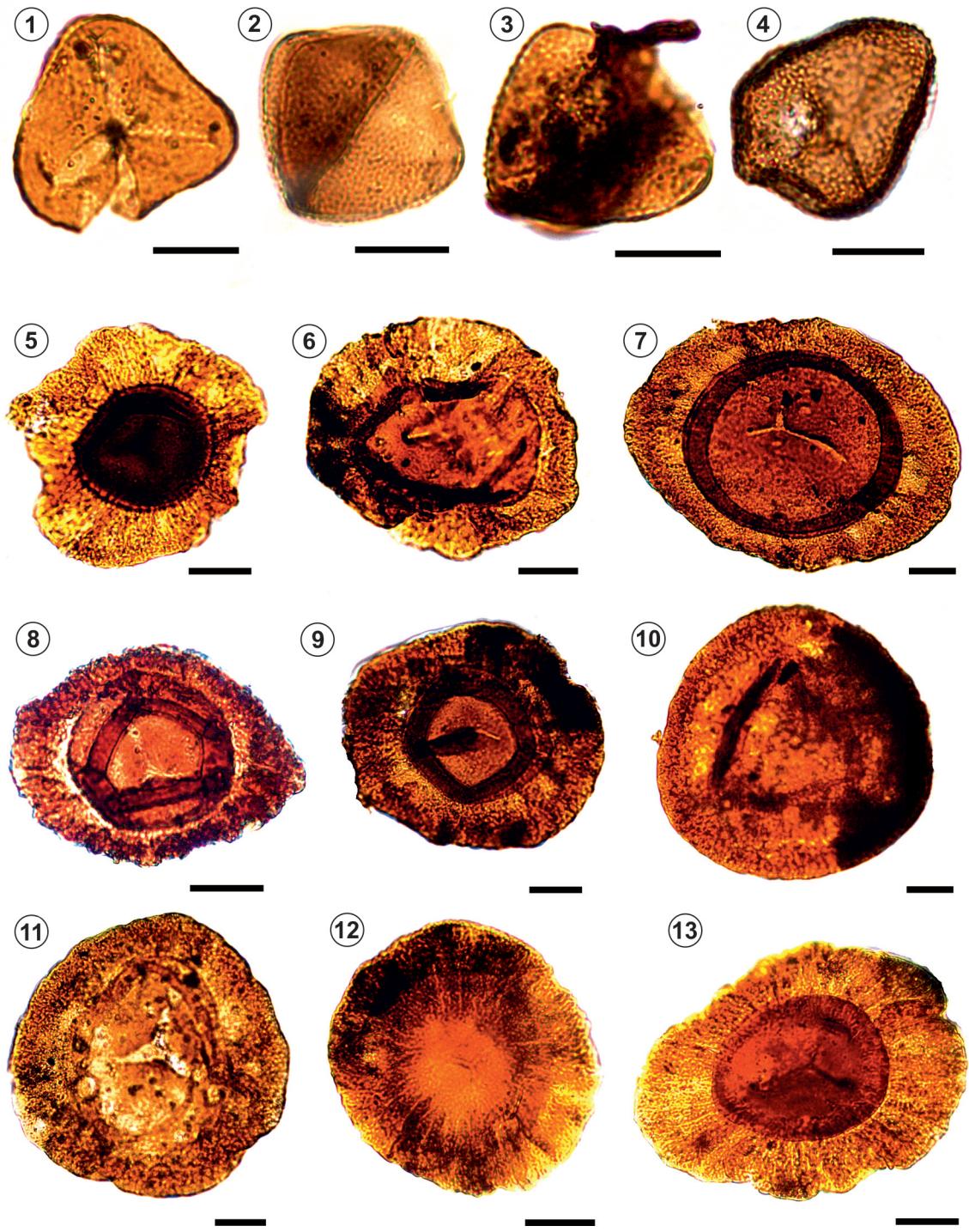


Figure 5. 1, *Granulatisporites austroamericanus* Archangelsky & Gamerro, 1979, BSIP 16629, EF: X.45.2; 2–3, *Pseudoreticulatispora confluens* (Archangelsky & Gamerro, 1979) Backhouse, 1991, 2, BSIP 16620, EF: K.51, 3, BSIP 16620, EF: X.34.1; 4, *Converrucosiporites micronodosus* (Balme & Hennelly, 1956) Playford & Dino, 2002, BSIP 16611, EF: Q.35.1; 5, *Cannanaropollis densus* (Lele, 1964) Bose & Maheshwari, 1968, BSIP 16598, EF: E.43.1; 6, 8–9, *Plicatipollenites gondwanensis* (Balme & Hennelly, 1956) Lele, 1964, 6, BSIP 16598, EF: N.52.1, 8, BSIP 16600, EF: P.48.1, 9, BSIP 16600, EF: W.34.1; 7, *Plicatipollenites malabarensis* (Potonié & Sah, 1960) Foster, 1979, BSIP 16599, EF: Q.30.4; 10–11, *Plicatipollenites trigonalis* Lele, 1964, 10, BSIP 16602, EF: F.11, 11, BSIP 16601, EF: E.48; 12, *Cannanaropollis janakii* Potonié & Sah, 1960, BSIP 16603, EF: G.64; 13, *Costatascyclus crenatus* (Felix & Burbridge, 1967) Urban, 1971, BSIP 16604, EF: W.46. Scale bars equal 20 µm.

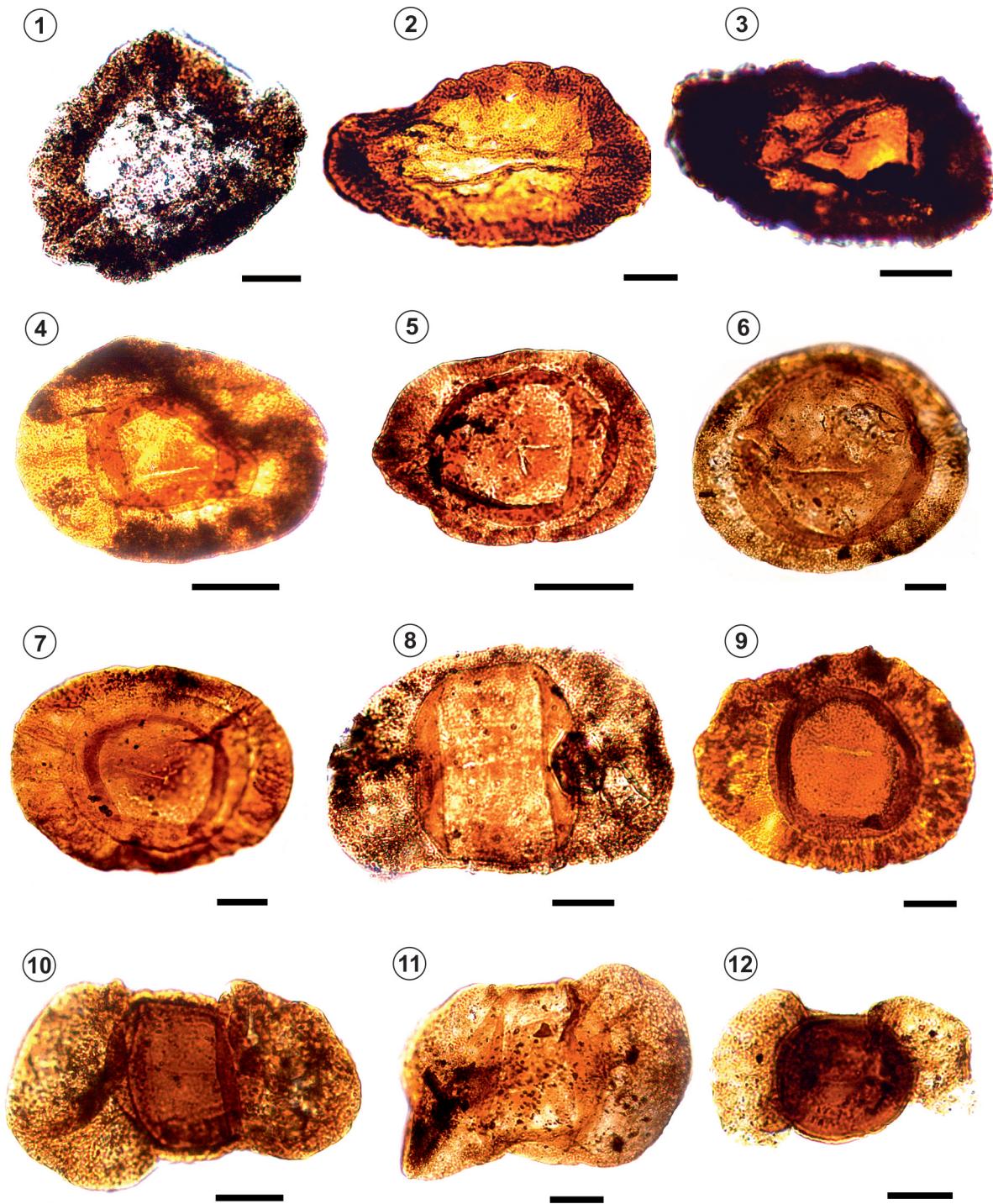


Figure 6. 1, *Cannanaropollis triangularis* (Mehta, 1944) Bose & Maheshwari, 1968, BSIP 16605, EF: X.51; 2, *Divarisaccus lelei* Venkatachala & Kar, 1966, BSIP 16603, EF: G.10; 3, *Latusipollenites quadrisaccatus* Marques-Toigo, 1974, BSIP 16608, EF: X.64; 4, *Potonieisporites densus* Maheshwari, 1967, BSIP 16609, EF: S.40; 5–6, *Potonieisporites novicus* (Bharadwaj, 1954) Poort & Veld, 1997, 5, BSIP 16600, EF: P.48.1, 6, BSIP 16609, EF: W.59; 7, *Potonieisporites congoensis* Bose & Maheshwari, 1968, BSIP 16609, EF: M.53.4; 8, *Potonieisporites barrelis* Tiwari, 1965, BSIP 16610, EF: L.71.4; 9, *Potonieisporites lelei* Maheshwari, 1967, BSIP 16602, EF: W.50.2; 10, *Potonieisporites lelei* transitional to *Limitisporites hexagonalis*, BSIP 16609, EF: T.61.1; 11, *Limitisporites hexagonalis* Bose & Maheshwari, 1968, BSIP 16609, EF: Y.61.1; 12, *Caheniasaccites flavatus* (Bose & Kar, 1966) Azcuy & di Pasquo, 2000, BSIP 16604, EF: U.34.2. Scale bars equal 20 µm (1–6, 9–12), 50 µm (7–8).

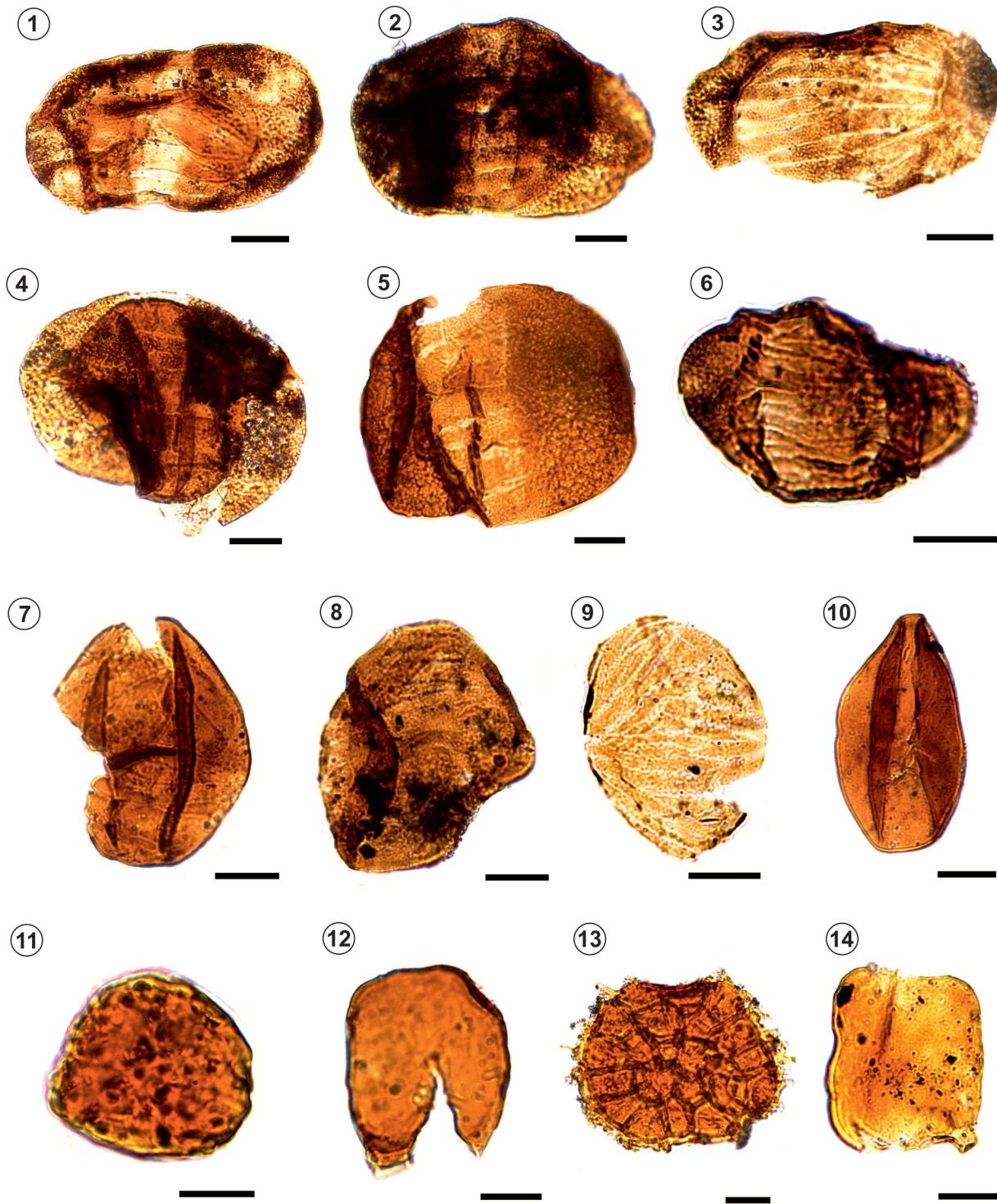


Figure 7. 1, *Illinites talchirensis* (Lele & Makada, 1972) Azcuy, di Pasquo, & Ampuero, 2002, BSIP 16604, EF: M.63.2; 2, *Protohaploxylinus perfectus* (Naumova, 1939) Samoilovich, 1953, BSIP 16611, EF: J.54; 3, *Vittatina subsaccata* Samoilovich, 1953, BSIP 16603, EF: S.69.3; 4, *Lunatisporites variesectus* Archangelsky & Gamerro, 1979, BSIP 16611, EF: J.54; 5, *Lunatisporites noviaulensis* Foster, 1979, BSIP 16602, EF: R.39; 6, *Striatoabietites multistriatus* Foster, 1979, BSIP 16613, EF: P.50.4; 7, *Pakahrites fusus* (Bose & Kar, 1966) Menéndez, 1971, BSIP 16614, EF: Q.47; 8, *Vittatina vittifera* (Luber in Luber & Waltz, 1941) Samoilovich, 1953, BSIP 16615, U.31; 9, *Marsupipollenites striatus* (Balme & Hennelly, 1956) Foster, 1975, BSIP 16616, EF: E.35.4; 10, *Marsupipollenites triradiatus* Balme & Hennelly, 1956, BSIP 16609, EF: Q.49.3; 11, *Botryococcus braunii* Kützing, 1849, BSIP 16626, EF: U.38.1; 12, *Brasilea scissa* (Balme & Hennelly, 1956) Foster, 1975, BSIP 16609, EF: S.32; 13, *Cymatosphaera gondwanensis* (Tiwari, 1965) Backhouse, 1991, BSIP 16601, EF: U.49; 14, *Tetraporina punctata* (Tiwari & Navale, 1968) Kar & Bose, 1976, BSIP 16599, EF: N.58.1. Scale bars equal 20 µm.

Figure 8. Age proposed for the Palynoassemblages I and II of the present study based on correlation with selected biostratigraphic schemes radiometrically constrained across Gondwana. References: Argentina-Paganzo Basin (Césari & Gutiérrez, 2001; di Pasquo *et al.*, 2010; Césari *et al.*, 2013; Césari & Chiesa, 2017; Valdez *et al.*, 2020); Bolivia, Tarija Basin (di Pasquo, 2003; di Pasquo *et al.*, 2017); Madre de Dios (di Pasquo *et al.*, 2015, 2019); Brazil, Paraná Basin (Beri *et al.*, 2012; Souza *et al.*, 2021); Uruguay, Paraná Basin (Beri *et al.*, 2011); Africa (Stephenson, 2009). Western Australia, Canning Basin (Foster & Waterhouse, 1988; Backhouse & Mory, 2020). Eastern Australia, Galilee Basin (Jones & Truswell, 1992; Backhouse & Mory, 2020). Abbreviations: AC, *Striatobalites anoverrucosus*-*Stauracerasites corduberensis*; BC, *Dictyostylites bireticulatus*-*Cristatisporites chacoparanensis*; FM, Fraile Muerto Member of Melo Formation; IS, *Cristatisporites incostans*-*Vittatina subsaccata*; KA, *Crassispora kosankei*-*Cystostyphus aczeyi*; MG, Mangruillo Member of Melo Formation; MR, *Converrucosporites micronodosus*-*Reticulatisporites reticulatus*; RS, Raistrickia *radiosa*-*Apicalisporis spinulistratus*; TB, *Marsupiolientes triradiatus*-*lunbladispora brasiliensis*.

TABLE 2. Biostratigraphic distribution of spores and pollen grain species across Gondwana addressed in the present study.

TABLE 2. Continuation.

Taxa	Basins	DAMODAR-RAJMAHAL (India)	PAGANZO (Argentina)	TARIJA (Bolivia)	MADRE DE DIOS (Bolivia)	PARANÁ (Brazil)	NAMIBIA (Uruguay)	CANNING (Western Australia)	GALILEE (Eastern Australia)											
	Samples	Pn	Pg	DMa	DMb	DMc	FS	KA-RS	BC	MR-TB	A1/A2	Vc	Ac	Cm	Vc	IS	Cc	Va-Mt	Pc	OpZ C OpZ D OpZ E
<i>Divariscoccus lelei</i>		X	X	X																
<i>Illinites talchirensis</i>			X	X	X															
<i>Latisporites quadrisaccatus</i>								X				X	X	X	X					X
<i>Limitisporites rectus</i>				X	X	X	X	X			X	X	X	X	X					X
<i>Limitisporites hexagonalis</i>			X	X	X	X	X				X	X	X	X	X					X
<i>Lunatisporites noviauensis</i>								X												X
<i>Lunatisporites varisectus</i>			X									X	X	X	X					X
<i>Marsupipollenites striatus</i>									X			X	X	X	X					X
<i>Marsupipollenites triradiatus</i>										X		X	X	X	X					X
<i>Parahipites fusus</i>							X					X	X	X	X					X
<i>Plicatipollenites gondwanensis</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Plicatipollenites malabarensis</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Plicatipollenites trigonalis</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Potoneisporites barrelis</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Potoneisporites congoensis</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Potoneisporites densus</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Potoneisporites lelei</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Potoneisporites magnus</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Potoneisporites neglectus</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Potoneisporites novicus</i>		X	X	X	X	X	X	X			X	X	X	X	X					X
<i>Protohaploxylinus latissimus</i>		X	X																	
<i>Protohaploxylinus limpidus</i>		X	X	X	X	X	X	X												
<i>Protohaploxylinus perfectus</i>																				
<i>Scheuringipollenites maximus</i>		X	X	X	X	X	X	X				X	X	X	X					
<i>Striatobaleites multistriatus</i>																				
<i>Vittatina subsaccata</i>												X	X	X	X					
<i>Vittatina vittifera</i>												X	X	X	X					

Abbreviations: Pn, Potoneisporites neglectus; Pg, Plicatipollenites negletus; DM, Parasaccites korbanensis; DM, Raistrickia densa-convolutispora muriornata; FS, Pakhapites fusu-s-Vittatina subsaccata; KA, Crassispora kosankei-Cystophytus azcuyi; RS, Raistrickia radiosa-Apicalisporites spinuliferatus; BC, Dicytorellites bireticulatus-Cristatosporites micronodosus-Reticulatisporites reticulatus; TB, Marsupipollenites triradiatus-Lundbladisporites brasiliensis; Vc, Vittatina costabilis; Mt, Ahrensporites cristatus; Cm, Cristatisporites inconspicua-Vittatina subsaccata; Cc, PC, Converrucosporites (= Pseudoreticulatisporites) confusus; Va, Vallatisporites arcuatus; Mt, Microbaculispora tentula. See also Figure 8. References: Damodar-Rajmatla basins (India); Tiwari and Tripathi (1992); Murthy et al. (2020); Paganzo Basin (Argentina); Césari and Gutiérrez (2001); di Pasquo et al. (2010); Césari and Chiesa (2017); Césari et al. (2013); Césari and Chiesa (2017); Valdez et al. (2020); Tarja Basin (Bolivia); di Pasquo (2003); di Pasquo et al. (2017); Madre de Dios (Bolivia); di Pasquo et al. (2015, 2019); Paraná Basin (Brazil); Mori et al. (2012); Souza et al. (2021); Paraná Basin (Uruguay); Beri et al. (2011); Namibia (Africa); Stephenson (2009); Canning Basin (Western Australia); Foster and Waterhouse (1998); Backhouse and Mory (2020); Galilee Basin (Eastern Australia); Jones and Truswell (1992); Backhouse and Mory (2020).

horizons and corresponding lithofacies, and associated fossils (see Murthy *et al.*, 2020, tab. 3) and proposed a correlation with radiometrically constrained similar assemblages across Gondwana. They established for the first time, a late Pennsylvanian age to the oldest *Potonieisporites neglectus* Assemblage Zone of Tiwari & Tripathi (1992) based on the absence of species with earliest Permian's FAD and the first occurrence of Glossopteridales in association with the overlying palynozones in the Damodar-Rajmahal Basins (Fig. 1.1). A Pennsylvanian age is also affirmed by global occurrence of the constituent elements of the palynozone in congruence with radiometrically constrained palynozones of South America and elsewhere in Gondwana (see di Pasquo *et al.*, 2015; Valdez *et al.*, 2020 and references therein).

The overlying *Plicatipollenites gondwanensis* Assemblage Zone from the siltstones above younger boulder beds was attributed to the early Cisuralian by Murthy *et al.* (2020) based on the first appearance of Permian marker taxa, reinforced by associated *Eurydesma* assemblage fauna (Lele & Makada, 1972). Currently, despite some taxonomic revision of species yet required, several palynozones across Gondwana, calibrated with radiometric data from interbedded igneous rocks, support or improve the stratigraphic ranges of Pennsylvanian/Permian diagnostic taxa that are present in the *P. gondwanensis* Zone (Fig. 8; Tab. 1).

Hitherto, Lele (1979) reported a palynoflora from the siltstone and clasts associated with the basal boulder beds of the Talchir Formation that unconformably overlie the Proterozoic Penganga limestones and dolomites near Irai in Maharashtra State, Central India (Wardha Basin, Fig. 1). He recorded for the first time a definite basal Talchir microflora dominated by phytoplanktonic species (70%), few trilete spores of *Punctatisporites* and monosaccate pollen species of the genera *Potonieisporites* (12%), *Cannanaropollis* (9.5%), *Plicatipollenites* (4.5%), *Caheniasaccites* (2.5%) and poorly preserved bisaccate pollen grains (striate?) (Tab. 1). Interestingly, while other researchers assigned this unit to the early Permian (e.g., Tiwari & Tripathi, 1992; Tiwari, 1999), Lele (1979) had correlated this assemblage with the late Pennsylvanian Stage 1 of Australia (Kemp *et al.*, 1977). Smith (1963) attributed a glacial origin to the Talchir boulder beds at Irai resting on striated pavements and pointed out its equivalence with other late Carboniferous Gondwanan glacial units.

The present study is the second palynofloral report of the Talchir Formation in the Wardha Basin. The older Palynoassemblage I (**PI**) of our study is tentatively correlated with that of Lele (1979) based on their compositional similarities despite the different kinds of phytoplankton in the present assemblage (Tab. 1). Lele (1979) did not find any equivalent assemblage with our Palynoassemblage II (**PII**). More accurate correlations between the two lower palynozones of the Talchir Formation across Gondwana are addressed below (Fig. 8).

Correlation with other basins in India

Despite the plethora of palynodata generated from the Indian Gondwana basins, unfortunately, most of the assemblages of the Talchir Formation cannot be correlated with our results and other Gondwanan palynoassemblages as they are mostly dated using dominance and sub-dominance of species instead of using chronometric tie species. Besides, some works do not provide a robust lithostratigraphic framework, which undermines the essence of biostratigraphic correlation. We have confined our correlation with the recent work of Murthy *et al.* (2020) from Rajmahal Basin, who applied chronometric tie species (FADs and chronostratigraphic ranges of species) in regional correlations of their assemblages. Hence, our Palynoassemblage I from the basal Talchir Formation correlates with the older Palynoassemblage I of Murthy *et al.* (2020) and the assemblage of Lele (1979), sharing palynological features like the dominance of monosaccate pollen species of the genera *Cannanaropollis*, *Plicatipollenites*, and *Potonieisporites* and the absence of typical Permian marker taxa. The Palynoassemblage II of the present study correlates with the younger assemblage from the Rajmahal Basin (Murthy *et al.*, 2020), characterized by more diversified pollen and spore species recording the first appearance of taxa such as *Cristatisporites pseudozonatus*, *C. conatus*, and *Lunatisporites variesectus*. However, the present assemblage documents the first appearance of several other taxa such as *Converrucosporites confluens*, *Latusipollenites quadrisaccatus*, *Marsupipollenites striatus*, *Pakhapites fusus*, *Striatoabieites multistriatus* (Tab. 2), documented in the latest Pennsylvanian–Asselian palynozones in Gondwana as further discussed.

Comparison and correlation with radiometrically constrained assemblages from Gondwana

Palynoassemblage I

Palynological associations documented across Gondwana since the Bashkirian are dominated by long-ranged monosaccate pollen grains with the absence of Permian markers like in the Palynoassemblage I of the present study. A comparison with radiometrically constrained palynoassemblages is addressed below to attempt correlations (Fig. 8; Tab. 2).

Argentina. The Carboniferous and Permian fossiliferous sequences of central-western Argentina (Paganzo, Río Blanco, Calingasta-Uspallata, and San Rafael Basins) contain abundant plant remains, palynomorphs, and invertebrates. The late Pennsylvanian biostratigraphic framework in Argentina referred to the *Raistrickia densa-Convolutispora muriornata* (DMZ) Assemblage Biozone is recognized in the Guandacol and Tupe formations of the Paganzo Basin and their stratigraphic equivalents, Jejenes, Lagares, and Agua Colorada formations, from different localities of San Juan and La Rioja provinces (Césari & Gutiérrez, 2001). Subzone A (DMA) of DMZ is characterized by the first record of monosaccate pollen genera *Plicatipollenites* and *Potonieisporites* and the absence of bisaccate striated pollen grains. The $^{206}\text{Pb}/^{238}\text{U}$ 319.57 ± 0.09 Ma and 318.79 ± 0.10 Ma ages of Gulbranson *et al.* (2010) from post-glacial transgressive facies of the Guandacol Formation was confirmed (320 ± 3 Ma) by Valdez *et al.* (2020), who added an age of 326 ± 3 Ma late Serpukhovian–Bashkirian to the Subzone A from basal glacial deposits (Pre-MTD1/Cycle 0 at Sierra de Maz).

Subzone B (DMB) of the DMZ is characterised by the presence of diverse monosaccate pollen and scarce taeniate pollen usually *Protohaploxylinus limpidus* (Césari, 1986; Césari & Gutiérrez, 2001; di Pasquo *et al.*, 2010). Subzone C (DMC) of the DMZ is similar in composition to DMB, composed of brackish and marine palynomorphs, especially acritarchs that constitute 70% of the assemblage, in association with monosaccate pollen grains lacking Permian markers (Gutiérrez & Limarino, 2006). Isotopic ages were obtained from different stratigraphic units in western Argentina, constraining these two zones mainly to the late Bashkirian/Moscovian–Kasimovian (Lech, 2002; Césari *et al.*, 2007, 2011, 2019; Gulbranson *et al.*, 2010). Despite PI shares only long-ranged monosaccate pollen grains with the DMC, the absence of Asselian diagnostic species that

appear in the conformably overlying PII (c. 4 m above) would support their correlation.

Brazil. The upper Paleozoic deposits of the Tubarão and Passa Dois groups in Paraná Basin of Brazil are well-represented by marine and continental environments with evidence of the interglacial and post-glacial events (Holz *et al.*, 2010). Souza (2006) proposed the Pennsylvanian *Ahrensisporites cristatus* Interval Zone (AcZ) for the basal portion of the Itararé Subgroup and the overlying *Crucisaccites monoletus* Interval Zone (CmZ) ranging approximately from the top of its lower portion to the middle portion of the unit. Souza *et al.* (2021) constrained these zones to the late Bashkirian to Kasimovian. Cingulizone, apiculate, and laevigate spores, and monosaccate pollen grains with radial and bilateral symmetry are dominant in both assemblages. Common constituents include *Cristatisporites* spp., *Vallatisporites* spp., *Punctatisporites gretensis*, *Calamospora* spp., *Apiculatisporis variornatus*, *Horriditriletes* spp., *Reticulatisporites* spp., *Spelaeotriletes ybertii*. Species of *Scheuringipollenites*, *Cycadopites*, *Limitisporites*, and *Protohaploxylinus* are also present in low frequencies (*i.e.*, lesser than 10% of the total assemblages). The *Vittatina costabilis* Interval Zone (VcZ) is mainly recorded from the upper Itararé Subgroup and the Rio Bonito Formation of the Passa Dois Group.

The lowermost glacial levels identified in the eastern area of the Brazilian Itararé Group is correlated to the *Ahrensisporites cristatus* Zone in the glacigenic rocks of the subsurface Lagoa Azul Formation (Souza, 2006; Holz *et al.*, 2010; Rosa *et al.*, 2019) and the DMA of the Guandacol Formation (Pérez Loinaze *et al.*, 2010; Valdez *et al.*, 2020). As in Argentina, Brazilian macrofloral species like *Nothorhacopteris argentinica* and *Botrychiopsis weissiana* characterized this interval with the absence of glossopeltids (Bernardes-de-Oliveira *et al.*, 2016). Despite the poorly diversified composition of our Palynoassemblage I, common species, along with its stratigraphic position close to the PII, allowed the correlation with the CmZ.

Bolivia. The Pennsylvanian–Permian Copacabana Formation in the Madre de Dios Basin is a biostratigraphically well-characterized succession (200–800 meters) dominated by fossiliferous carbonates and intercalated siliciclastic facies in some intervals. Thin ash beds along this unit are numerous in different localities and in the core from the Mobil-Oxy Manuripi X-1 exploration well in northern Bolivia nearby

Pando X-1 cores (di Pasquo *et al.*, 2019 and references therein). U-Pb zircon age determination carried out in fractured siliceous green tuffaceous interval at depths of 882.4–883.2 meters yielded the average $^{206}\text{Pb}/\text{U}_{238}$ CA-ID-TIMS age of 316.0 ± 0.4 Ma supporting a Bashkirian through Middle Moscovian age for calcareous foraminifera, conodonts, and palynomorphs (A1 and A2 palynofloras in Table 2) documented (Hamilton *et al.*, 2016; di Pasquo *et al.*, 2016, 2019). In southern Bolivia and northern Argentina, the Pennsylvanian *K. volkheimeri*-*C. azcuyi* (Kv-Ca) Superzone, divided into five First Appearance Interval Zones (Fig. 8), is documented in the Macharetí and Mandiyutí groups of the Tarija Basin (di Pasquo, 2003; di Pasquo *et al.*, 2017 and references therein). Long-ranged *Punctatisporites gretensis* and monosaccate pollen grains are shared between PI and the Pennsylvanian assemblages of Bolivia, and along with the stratigraphic position of the PI, a correlation with the MR-TB Zones is established.

Australia. Palynological studies in the Eastern Australian Joe Joe Group of the Galilee Basin by Jones & Truswell (1992) allowed the characterization of the Pennsylvanian (Bashkirian–Kasimovian) *Spelaeotriletes queenslandensis* (= *Grandispora queenslandensis* sensu Playford *et al.*, 2001) Superzone (equivalent to the *Spelaeotriletes ybertii* Assemblage of Powis 1984), subdivided into three Oppel-zones A–C followed by the Oppel-zones D (Kasimovian–Asselian) and E (Asselian). The basal Oppel-zone A (OP A) contains monosaccate pollen, without taeniate specimens, like the Subzone A of the Argentinian DM Zone and the correlative zones of South America depicted in Figure 8. Fielding *et al.* (2008) correlated the OP A with the glacial interval C2 of Australia. Roberts *et al.* (1995) reported for the *Levipustula levis* marine fauna a radiometric age of 321–323 Ma (SHRIMP zircon dating). This agrees with the inferred age of the Guandacol Formation glacial event of the Argentinian and coeval glacial-marine deposits (*i.e.*, Hoyada Verde Formation) bearing rich invertebrate assemblages mainly belonging to the *Levipustula* Zone (Césari *et al.*, 2007; Vergel *et al.*, 2015). The first striated pollen grains of the genus *Protohaploxylinus* along with *Cristatisporites pseudozonatus* are documented in the OP C, and *Horriditriletes ramosus* appears in the OP D, all persisting up to the OP E. These species recorded in our PII make a correlation of PI with any of these zones difficult to assert.

Palynoassemblage II

The qualitative composition of Palynoassemblage II compared with other radiometrically constrained assemblages to the Gzhelian–Asselian (Permo/Carboniferous boundary) across Gondwana is addressed, and correlations proposed (Fig. 8; Tab. 2).

Argentina. The *Pakhapites fusus*-*Vittatina subsaccata* (FS) palynological Zone defined by Césari and Gutiérrez (2001) was defined in the stratigraphic units originally referred to the Asselian–Sakmarian in western Argentina. The appearance of *Gangamopteris* megafloristic Zone coincides with the base of the FS Zone. An oldest age of 298–301 Ma dated the first appearance of glossopterid remains that appeared interbedded with basaltic horizons dated between K/Ar 293 ± 6 Ma and 308 ± 6 from the lower part of the La Colina Formation (Thompson & Mitchell, 1972; Limarino & Césari, 1984). Gutiérrez and Limarino (2006) proposed the Río del Peñón Formation at the Río Blanco area in northern Precordillera, as a potential stratotype for the boundary between the DM and FS Zones. Gulbranson *et al.* (2010) reported a $^{206}\text{Pb}/\text{U}_{238}$ age of 310.63 ± 0.1 Ma obtained from fluvio-deltaic deposits of the middle part of this unit below the FS Zone. An increase of striate pollen grains, the first appearance of *Converrucosporites confluens*, and species of *Vittatina* and *Pakhapites fusus*, *Latusipollenites quadrisaccatus*, *Marsupipollenites striatus*, *Striatoabieites multistriatus* characterized this zone (see also di Pasquo *et al.*, 2010; Césari *et al.*, 2013; Césari & Chiesa, 2017). These species shared with our Palynoassemblage II supporting their correlation (Fig. 8; Tab. 2).

Brazil and Uruguay. The *Crucisaccites monoletus* Interval Zone (CmZ) of Souza (2006) ranges approximately from the top of the lower portion to the middle portion of the Itararé Subgroup in the Paraná Basin of Brazil. The upper limit characterized by the appearance of *Illinites unicus* and the first species of the genus *Vittatina* marks the beginning of the VcZ. The VcZ is radiometrically constrained based on ID-TIMS U-Pb zircon ages ranging from 299 ± 2.6 to 296 ± 1.4 Ma from ton steins located in the middle section of the Rio Bonito Formation (see Mori *et al.*, 2012; Griffis *et al.*, 2018; Valdez *et al.*, 2019, 2020). Stephenson (2009) pointed out that the base of the *Vittatina costabilis* Interval Zone is older than the age of the base of the Rio Bonito Formation and may be approximated to the age of the first occurrence of

Converrucosporites confluens in the Carboniferous/Permian boundary of Namibia (*i.e.*, 298.9 ± 0.31 – -0.15 Ma; Ramezani *et al.*, 2007). Recently, Souza *et al.* (2021) verified that VcZ is positioned between the Gzhelian and the Artinskian, based on radiometric ages obtained from levels containing spore-pollen associations assigned to this zone.

The San Gregorio up to the lower part of the Melo Formations (Frayle Muerto Member) in Uruguay are referred to the *Cristatisporites inconstans-Vittatina saccata* Assemblage Zone (IS) characterized by trilete spores and monosaccate pollen and, to a lesser extent by non-taeniate bisaccate, taeniate bisaccate, and plicate pollen grains (Beri *et al.*, 2011). The overlying *Striatoabieites anaverrucosus-Staurosaccites cordubensis* (AC) Assemblage Zone is constrained by radiometric datings from ashfall deposits of the Mangrullo Member of the Melo Formation that yielded U-Pb SHRIMP ages of 275.9 ± 4.8 Ma and 269.8 ± 4.7 Ma (Rocha-Campos *et al.*, in Beri *et al.*, 2011). Most the diagnostic species of Vc and AC Zones, such as *Converrucosporites confluens*, *Pakhapites fusus*, *Latusipollenites quadrisaccatus*, *Marsupipollenites striatus*, *Striatoabieites multistriatus*, *Vittatina subsaccata*, among others, occurred in our Palynoassemblage II, supporting their correlation (Fig. 8; Tab. 2).

Bolivia. Two Cisuralian palynoassemblages were recorded in marine and transitional deposits of the Copacabana Formation at Apillapampa in central Bolivia (di Pasquo & Grader, 2012). First appearances of mainly cosmopolitan diagnostic taxa allowed the characterization of a lower assemblage assigned to the *Vittatina costabilis* Zone of Paraná Basin and the upper one that occurred in overlying marine and coal-bearing transitional intervals linked to the *Lueckisporites virkkiae* Zone. U-Pb geochronology (ID-TIMS of zircons) from the lowermost tuff sample of five interbedded tuffs analyzed provided an absolute age of 298.9 ± 0.15 Ma (earliest Asselian) and the uppermost dated late Sakmarian ($291.6 \text{ Ma} \pm 0.9$ Ma). This interval confirmed previous microfossil (conodonts, fusulinids) ages and supports that the plant remains belonging to the *Glossopteris* Flora recorded in the Coal Member appeared in the late Sakmarian (di Pasquo *et al.*, 2015). A correlation is proposed herein between the present PII of the Talchir Formation and the lower VcZ (Fig. 8) based on their common pollen elements such as *Vittatina subsaccata*, *V. vittifera*, *Pakhapites fusus*, *Marsupipollenites striatus*, *Lunatisporites variesectus*, and several species of

monosaccate pollen genera *Cannanoropolis*, *Potonieisporites*, *Plicatipollenites* among other taxa (Tab. 2).

Australia. The Palynoassemblage II of the present study shares some index taxa with the *Converrucosporites confluens* Oppel Zone (=*Pseudoreticulatispora confluens*) recognized by Foster and Waterhouse (1988) in core intervals from the Calytrix No. 1 Borehole and in several boreholes by Backhouse and Mory (2020) in the Canning Basin of Australia. Those species are *Brevitriteles cornutus*, *C. confluens*, *Cycadopites cymbatus*, *Horriditriteles ramosus*, *Cristatisporites pseudozonatus*, *Marsupipollenites striatus*, *Protohaploxylinus limpidus*, *Striatoabieites multistriatus*, and monosaccate pollen grains. Instead, the species of *Vittatina* found in our PII were not in the Canning Basin (Backhouse & Mory, 2020).

Africa. The Palynoassemblage II shared with the *Converrucosporites* (=*Pseudoreticulatispora*) *confluens* assemblage from Carboniferous–Permian Ganigobis shale Member (Dwyka Group) in Namibia (Stephenson, 2009) species such as *Horriditriteles ramosus*, *H. uruguaiensis*, *Cycadopites cymbatus*, *Brevitriteles cornutus*, *Vittatina*, and *Converrucosporites confluens*. In this succession, Bangert *et al.* (1999) identified three sets of ash layers (termed I, II, and III). The upper ash layer of Set II (Ash Layer IIb at $25^\circ 53' 35''$ S, $18^\circ 00' 51''$ E, south of Ganigobis) gave a SHRIMP-based age from juvenile magmatic zircons of 302.0 ± 3.0 Ma (Bangert, 2000) so placed it entirely within the Pennsylvanian, Gzhelian or Kasimovian. Bangert *et al.* (1999) also provided an age for other ash layers within the Ganigobis Shale Member of 299.2 ± 3.2 Ma from a road cutting east of Ganigobis, near Tses. Both ages are still substantially older than the age corresponding to the lower limit of the range of the *Converrucosporites confluens* Oppel Zone in Western Australia (approximately 295 Ma). Stephenson (2009) proposed that the age of the *Converrucosporites confluens* Oppel Zone in Western Australia may range earlier than the mid–late Asselian previously established considering the radiometric dates from the Ganigobis Shale Member.

Barbolini *et al.* (2018) carried out a comprehensive review of previous palynological works in South Africa and identified biostratigraphically-restricted palynomorphs from the western, southern, and northeastern facies of the upper Dwyka (Asselian–Sakmarian K0 Zone) and Ecca Group (Artinskian–Lopingian K1–K11 Zones) in the Main

Karoo Basin. The chronostratigraphic framework of the Karoo successions was supported by isotopic dates (Barbolini *et al.*, 2016). Several species are shared with the K1Z like the possible pan-Gondwanan Carboniferous–Permian index taxa *Pakhapites fusus* and *Vittatina*, among other spores (*Brevitriletes cornutus*, *Con verrucosporites micronodosus*, *Horriditriletes ramosus*, *Verrucosporites andersonii*) and pollen grains (*Cannanoropollis methae*, *Cycadopites cymbatus*, *Marsupipollenites striatus*, *Plicatipollenites gondwanensis*, *Protohaploxylinus limpidus*, *Striatoabieites multistriatus*). A correlation of our PII with the K1Z is impeded due to their significant age difference.

CONCLUSIONS

The present work is significant because it is the second palynologic study of the Talchir Formation from Wardha Basin in India. Two palynoassemblages characterized by successive appearance of species and their biochron were compared and correlated with palynozones constrained by independent faunal dates and radiometric dates obtained across Gondwana. Quantitative abundance of genera is the conventional practice in Indian palynostratigraphic studies that may or may not be chronologically significant worldwide as it may reflect responses to specialized ecological conditions. Despite the absence of radiometric dates from the Indian Gondwana sequences, concerted efforts are being made to provide absolute dates in the near future. Meanwhile, Palynoassemblage I from the basal-most samples of the Borehole 131 from Penganga area is assigned to the oldest *Potonieisporites neglectus* Zone of Tiwari & Tripathi (1992) from the Indian Gondwana Sequence. The Palynoassemblage I composed of monosaccate pollen grains and few trilete spores long-ranged from Pennsylvanian to Permian is likely akin to the late Pennsylvanian by correlation to the radiometrically constrained late Serpukhovian–late Pennsylvanian DM Zone of central-western Argentina.

The Palynoassemblage II is constrained to the Carboniferous/Permian boundary by correlation with *Pakhapites fusus*-*Vittatina subsaccata* Zone of central-western Argentina, based on common diagnostic species of *Vittatina* and *Pakhapites* and *Con verrucosporites confluens*, among others. The relative age of the base of this zone is accurately assigned to the Kasimovian/Gzhelian based on $^{206}\text{Pb}/^{238}\text{U}$ age of 310.63 ± 0.1 Ma obtained from the middle

part of the Rio del Peñón Formation at the Rincón Blanco section in Mendoza Province of Argentina.

The lower limit is also consistent with the associated marine fauna correlatable with the *Tivertonia-Streptorhynchus* Zone of Sabattini *et al.* (1991) and the DMb and DMc of Tupe Formation (Paganzo Basin, San Juan Province, Argentina), confirmed by radiometric dates. The *Potonieisporites neglectus* Zone of Tiwari & Tripathi (1992) was attributed to the earliest Permian and very few works to the latest Pennsylvanian. We suggest a constraint to the late Pennsylvanian (Kasimovian–Gzhelian) considering the above biostratigraphic correlation together with the stratigraphic position of the studied lower interval of the Talchir Formation lying directly over Precambrian rocks. A Carboniferous–Permian age is also confirmed by correlation with the *Con verrucosporites confluens* assemblage from the Ganigobis lower shale Member (Dwyka Group) of Namibia, South Africa. The isotopic data obtained from the interbedded ash layers yielded ages of 299.2 ± 3.2 Ma and 302.0 ± 3.0 Ma, extending into the late Asselian–Sakmarian.

The integration of our palynostratigraphic information from the Talchir Formation with palynologic, radiometric and faunal dates is applied herein to propose a more accurate age for the two palynoassemblages documented and their correlation with late Kasimovian–Gzhelian to Asselian palynozones of southern South America, Africa, Australia and correlative palynofloras of this age interval elsewhere in Gondwana. Future work should focus on providing radiometric data to improve current palynostratigraphic knowledge and known biozonation schemes in a robust lithostratigraphic framework from the Indian Gondwana basins. In this context, the Talchir Formation should be our main target as it was overlooked in favor of the economically important coal-bearing Barakar and Raniganj formations.

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